

STRATOSPHERIC AEROSOL CONTENT AND VARIABILITY: BACKGROUND CONDITIONS and CONTROL BY VOLCANIC ERUPTIONS AND WILDFIRES

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¹²National Institute of Aerospace (NIA), Hampton, VA, USA

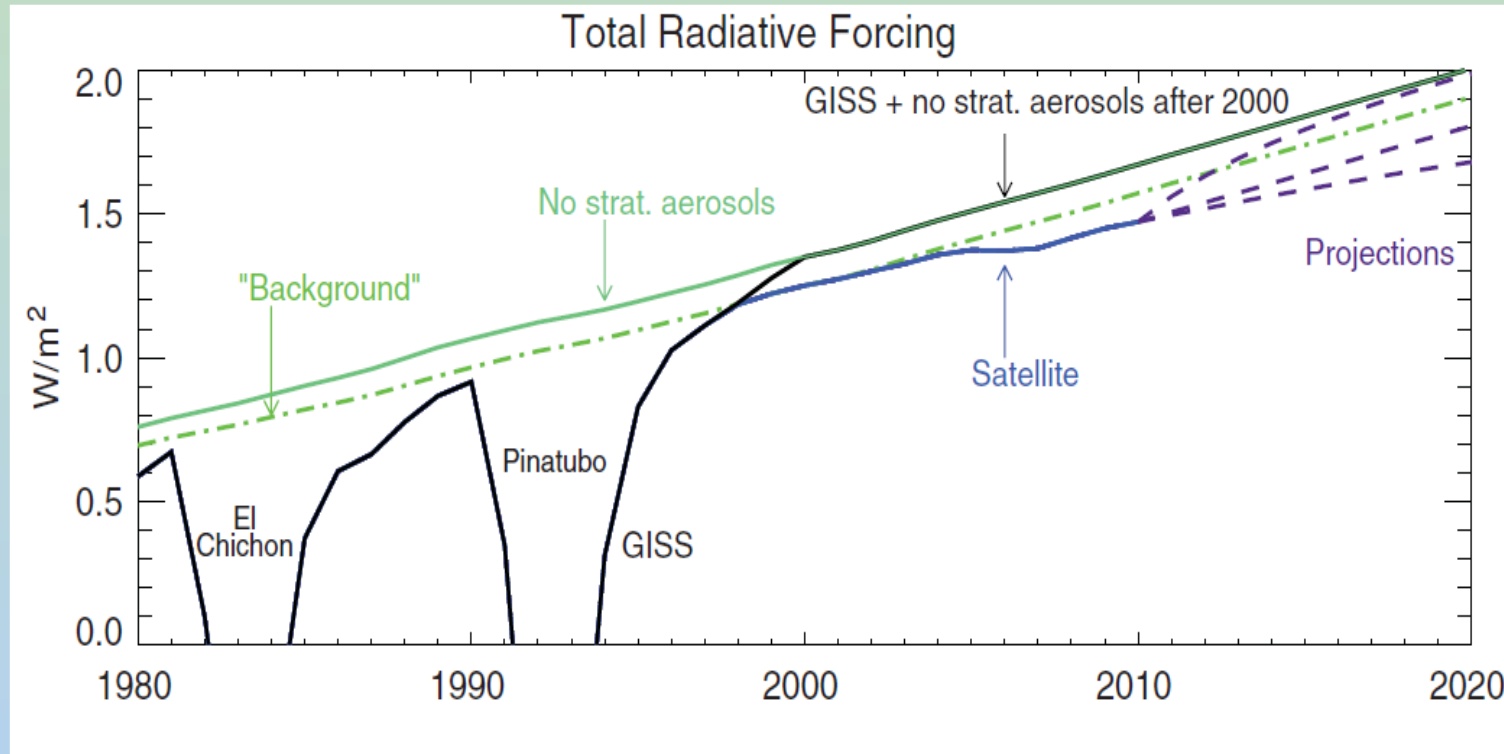
¹³Federal Institute of Technology (ETHZ), Zurich, Switzerland

HEMERA Workshop, June 5th 2022



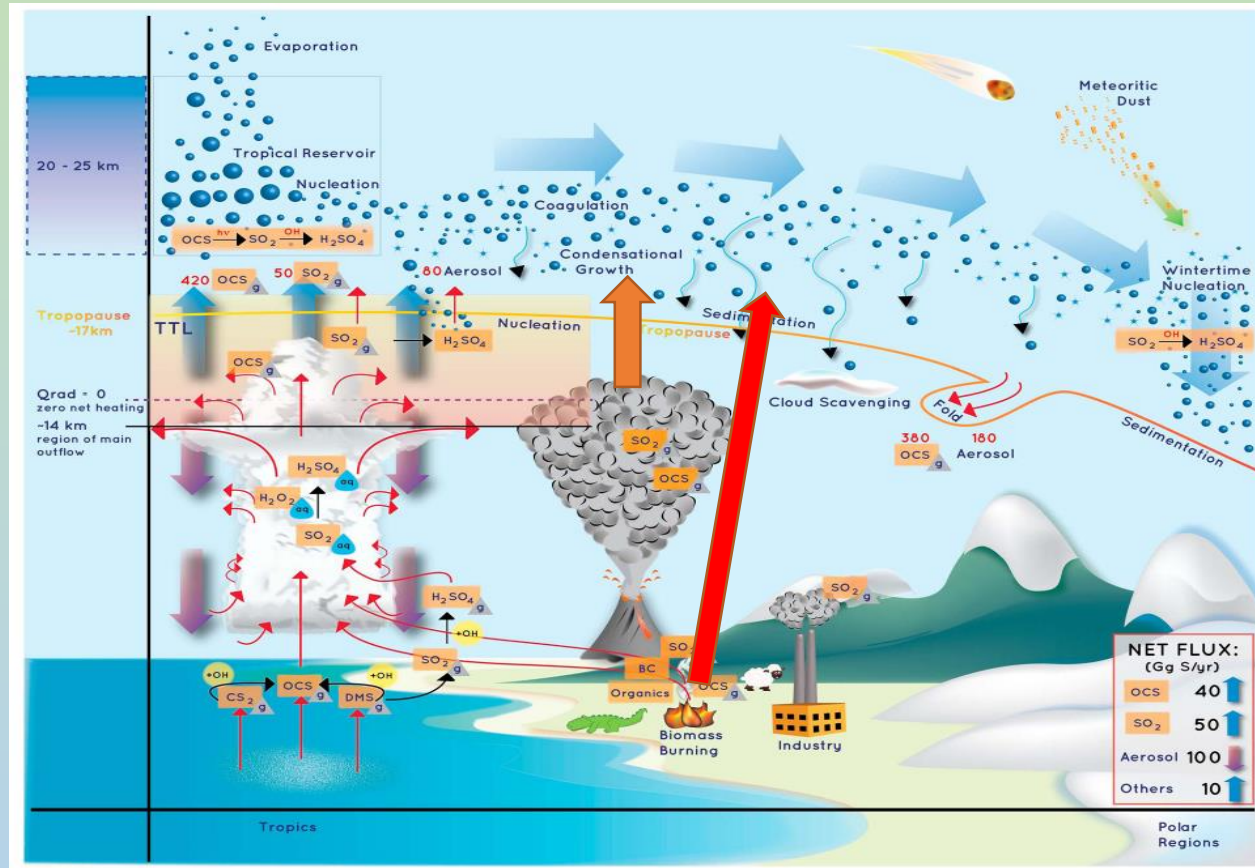
Radiative impact of stratospheric aerosols

Solomon et al., Science, 2011



-0.07°C due to stratospheric aerosols

Cycle of stratospheric aerosols



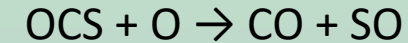
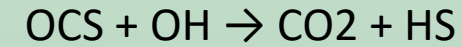
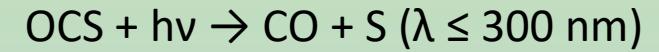
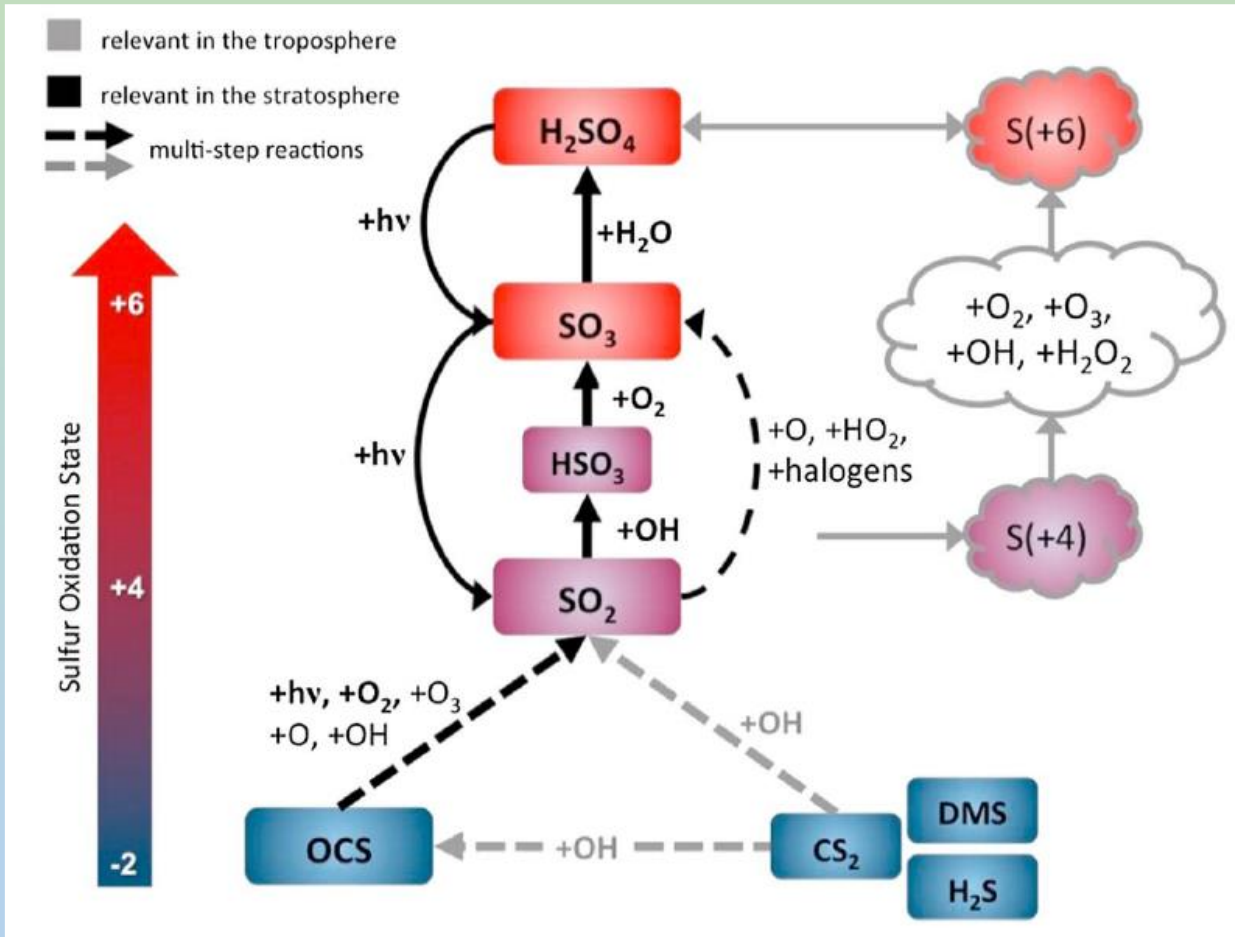
Kremser et al., Rev. of Geophys., 2016

Photolysis + oxydation of OCS in the stratosphere (contribution=56% in Sheng et al. JGR (2015) → emitted in the troposphere mainly from oceans, biomass burning (10-20%), wet areas, and anthropogenic activities; Bruhl et al. ACP

Oxydation of SO₂ in the stratosphere (contribution=44% in Sheng et al. JGR (2015) → anthropogenic activities and natural emissions (DMS, H₂S, volcanic degassing) with a similar contribution

Contribution of SO₂ from Chinese coal not significant (Neely et al. GRL, 2013; Sheng et al. JGR, 2015)

Sulfur oxidation



Very important role of sulfate aerosols on stratospheric ozone chemistry



=> Reduction of NO_x and of their ozone loss efficiency in the middle stratosphere

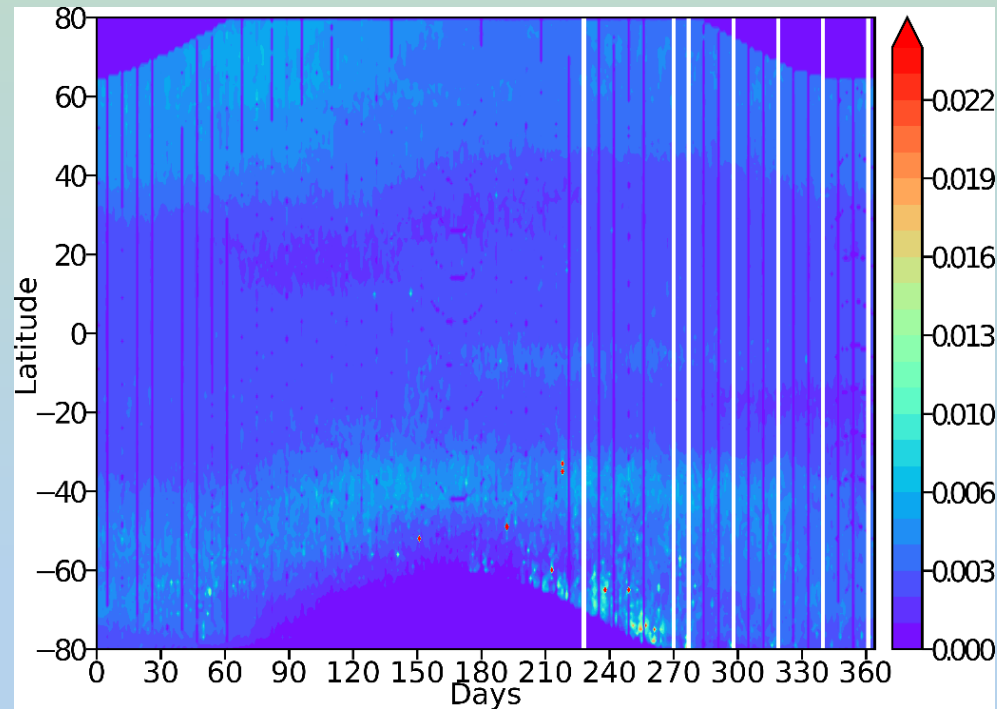
=> Increase of ozone loss by ClO_x, BrO_x et HO_x in the lower stratosphere

Kremser et al., Rev. of Geophys., 2016

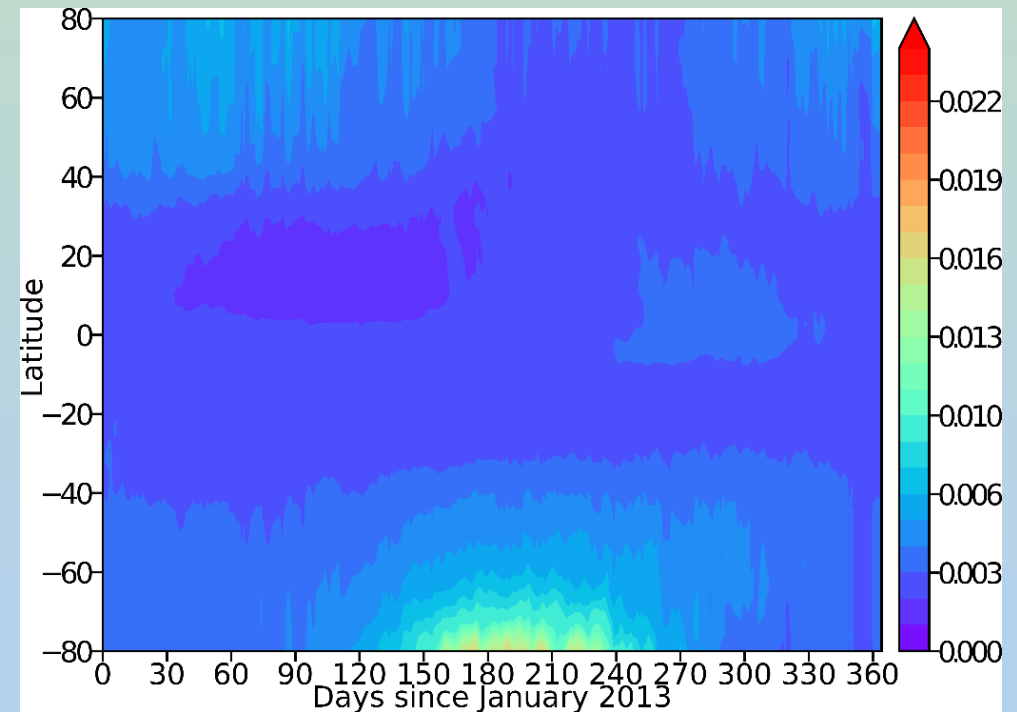
« Background » stratospheric aerosol

Latitudinal distribution of Aerosol Optical Depth (AOD)

OMPS, year 2013
675 nm

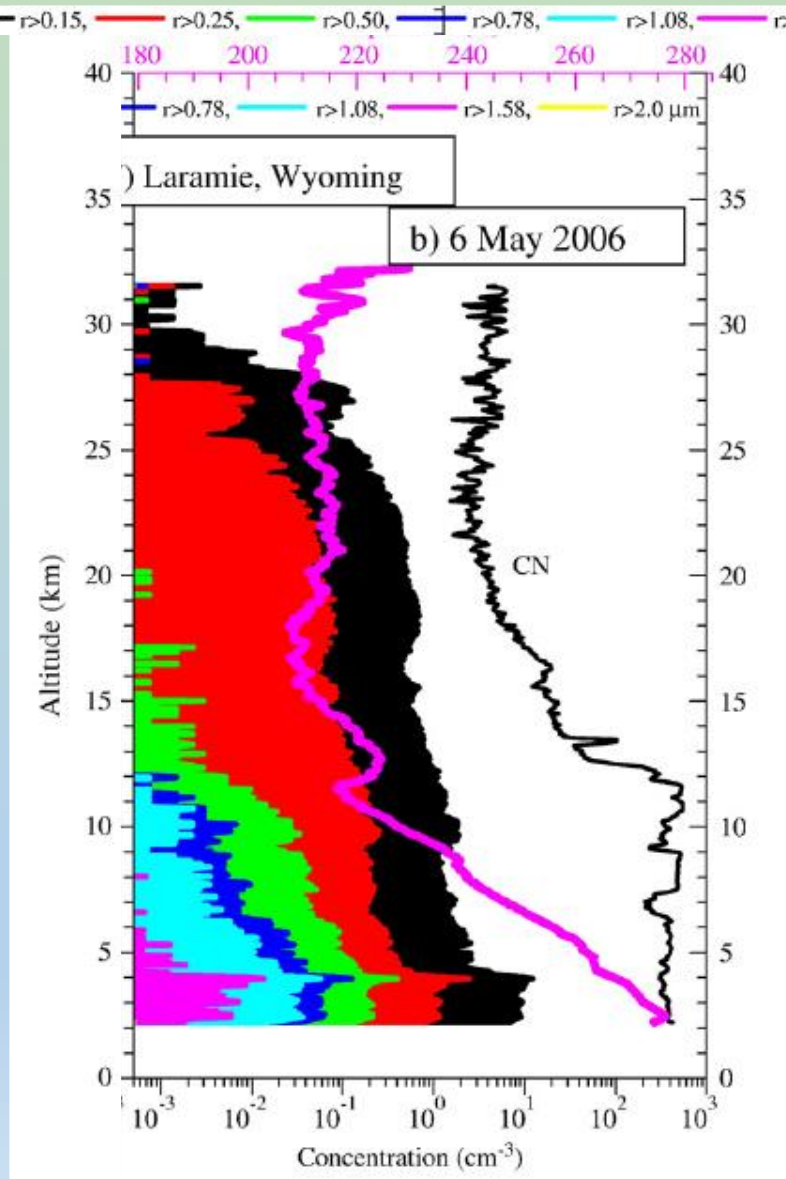


CESM1(WACCM)-CARMA, Year 2013



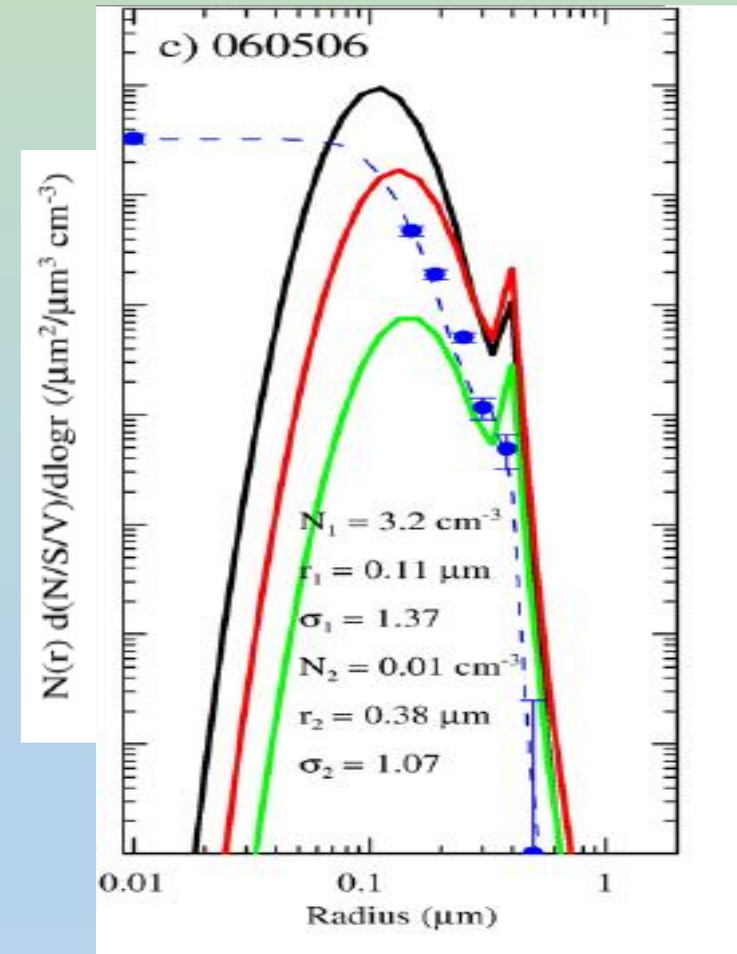
-> Slight seasonal cycle and latitudinal variations (tropopause temperatures and height, gas precursors, polar vortex)

« Background » stratospheric aerosol population from in situ observations

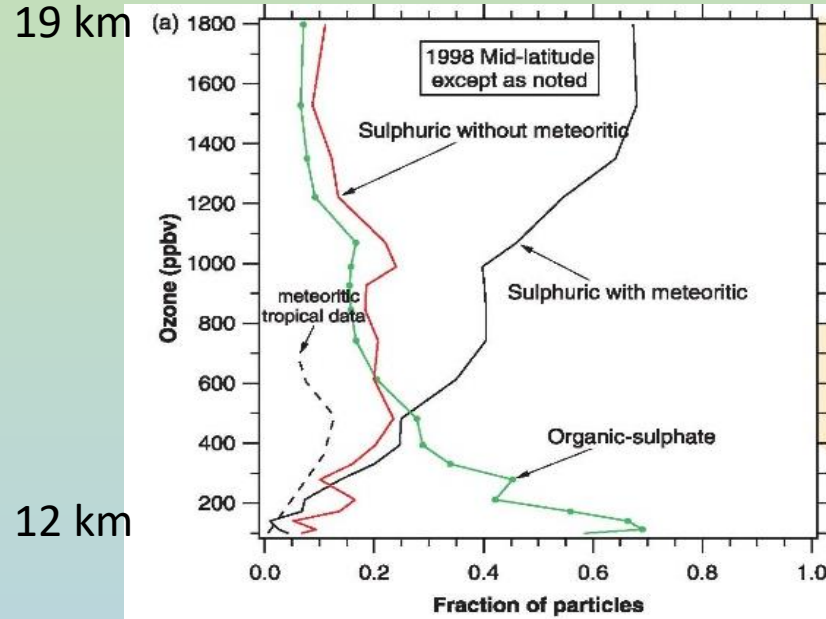


Legend for Figure c):

- dndlogr (black line)
- dsdlogr (red line)
- dvdlogr (green line)
- Integral ($dN/dr dr$) (dashed blue line)
- Measurements (blue circles)



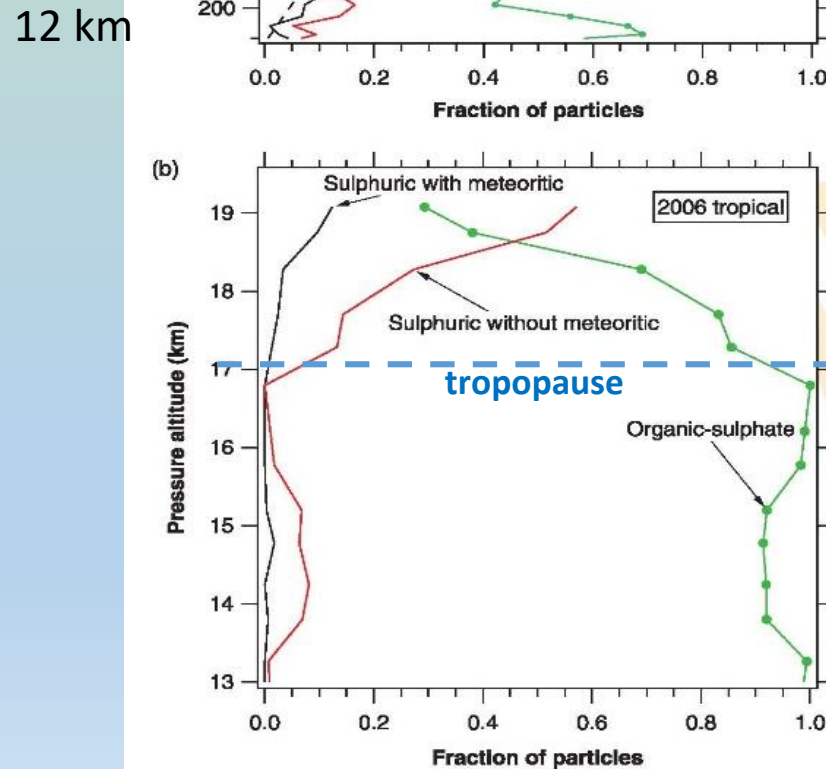
Particle composition in the lower stratosphere: mass spectrometry observations from aircraft



Mid-latitudes

Murphy et al., QJRMS, 2014

Sulfuric acid particles with meteoritical material dominate above ~15 km



Tropical-Latitudes

Mixture between sulfate and organic particles largely dominate in the tropical UTLS

Black carbon in the lower stratosphere

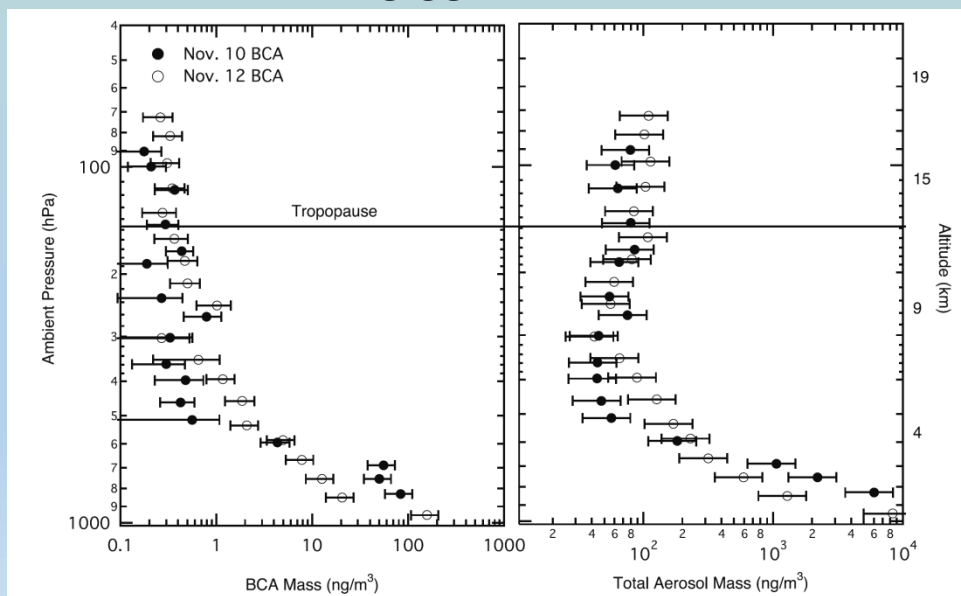
In fire plume quiescent period (or unreported events)

▪ *Strawa et al., JGR, 1999*

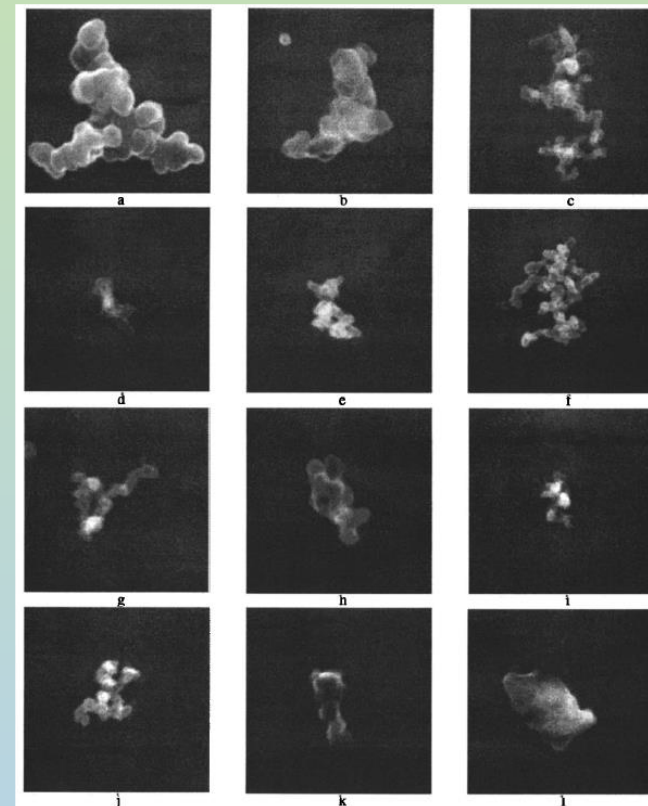
Wire impactors

Black Carbon aerosol number densities ~1% of the total number

28-38°N



19 km 50°N



95% in the shape of aggregates

▪ *Schwarz et al., JGR, 2006*

- Soot photometer: laser-induced incandescence to detect refractory particles 0,15 - < 1 μ m
- Black carbon aerosols: <1% of total aerosol mass in the stratosphere
- at least 40% have “coating” (= internally mixed)

Refractory particulates in the lower stratosphere in the polar vortex (1/2)

Impactor for aerosol collection (M55 Geophysica)
 + laboratory transmission and scanning electron spectroscopy (TEM and SEM) for size/morphology and by EDX (Energy-dispersive X-ray) for composition.



Schutze et al., ACP, 2017

Aerosol collection in the Arctic stratosphere ER2 aircraft, Jan-Feb 2000
Mainly carbonaceous particles associated with metals

Composition:

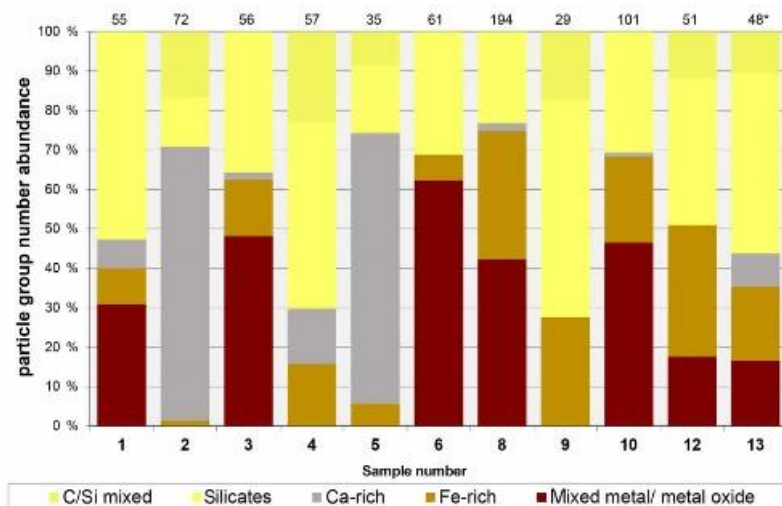
- Particles with different morphologies than Ebert et al.'s results
- Only **amorphous refractory carbonaceous particles** collected (among the sulfate population) C+O (72-100%) + others **with metals**.
- Refractory carbonaceous particles not included in or coated by sulfate are surprisingly dominant
 -> in contradiction with results from Mass Spectrometry.
- Concentrations : a few percent of the total number of particles in the air (between 0.62 and 2.3 (mg.air)⁻¹)

Ebert et al., ACP, 2016

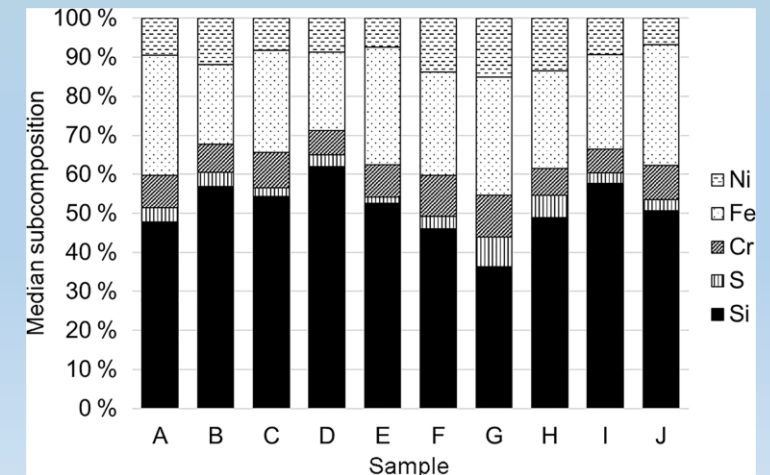
End of winter 2010: more refractory particles from the mesosphere
 - sizes >0,5 μm : **silicate spheroids, silicate/carbon mixture, iron-rich particles, calcium-rich particles, complex mixture of metals**
 - sizes <0,5 μm : **Presence of soot**
 - estimated concentrations : 10⁻² cm⁻³ for sizes of 0.75 μm

Arctic stratosphere, winter 2010 (Ebert et al., ACP, 2016)

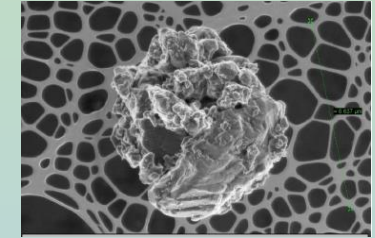
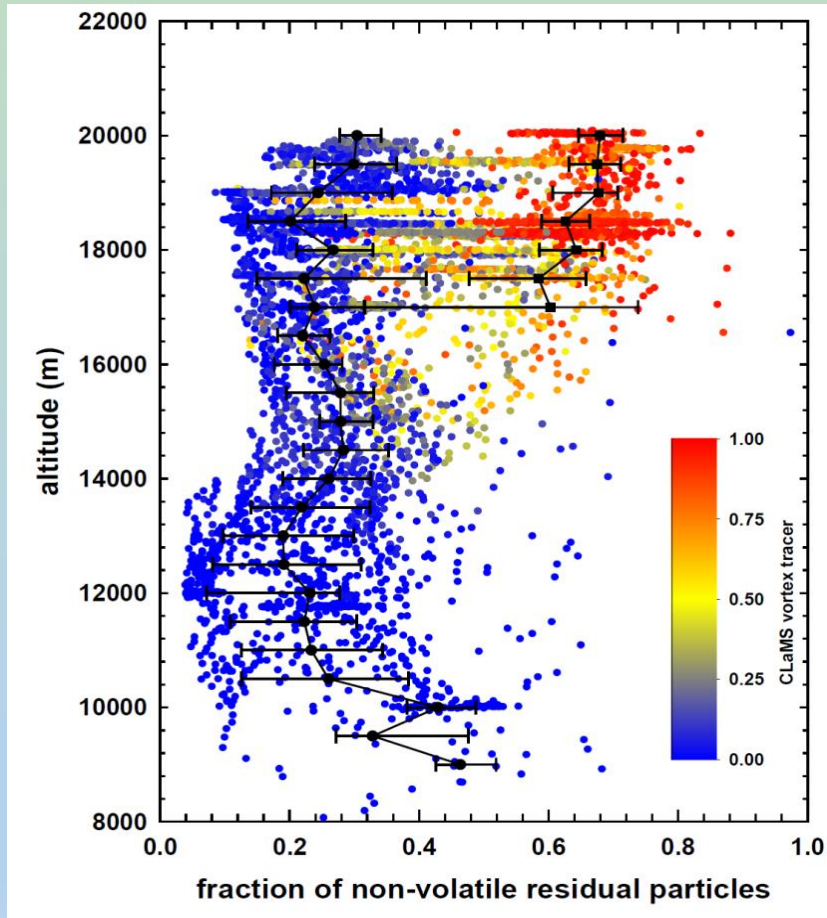
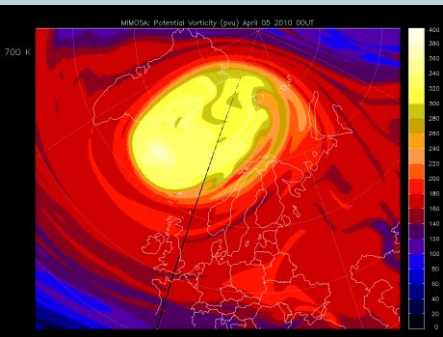
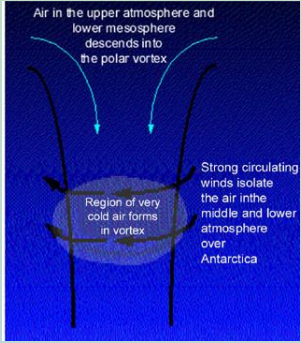
Up to 21 km



Arctic stratosphere, Jan-Feb 2000 (Schutze et al., ACP, 2017)



Refractory particulates in the lower stratosphere in the polar vortex (2/2)



February-March 2003

Aerosol counter with particle vaporisation

Size range 0.4-23 μm

In the polar vortex:

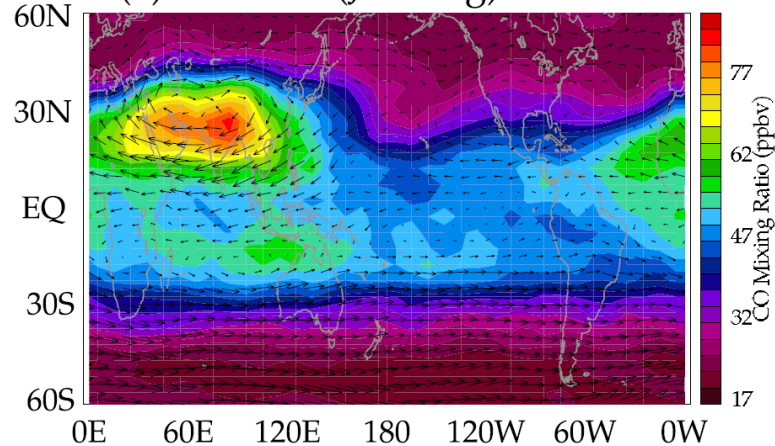
→ more than 60% contain non-volatile particles / Total concentration

→ Meteoric material transported from the mesosphere

Asian monsoon anticyclone: presence of pollutants and aerosols in the tropopause region (1/4)

~15-19 km altitude range

(a) MLS CO (Jul-Aug) 100 hPa



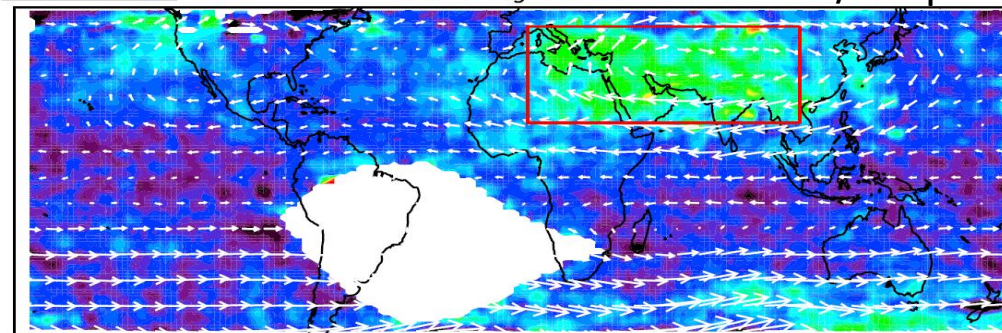
Park et al. (JGR, 2006)

Asian Tropopause Aerosol Layer (ATAL):

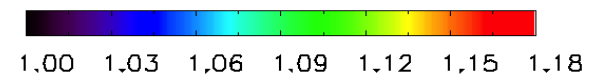
100 hPa

Jul-Aug 2006

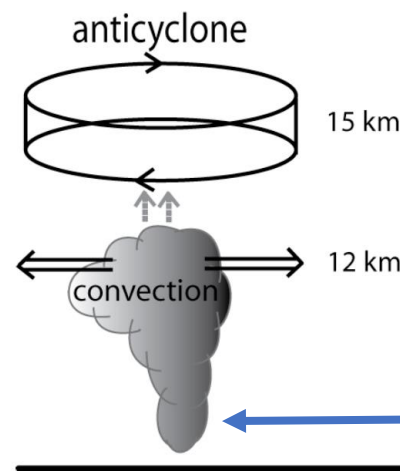
CALIOP/Calipso



SR@532nm

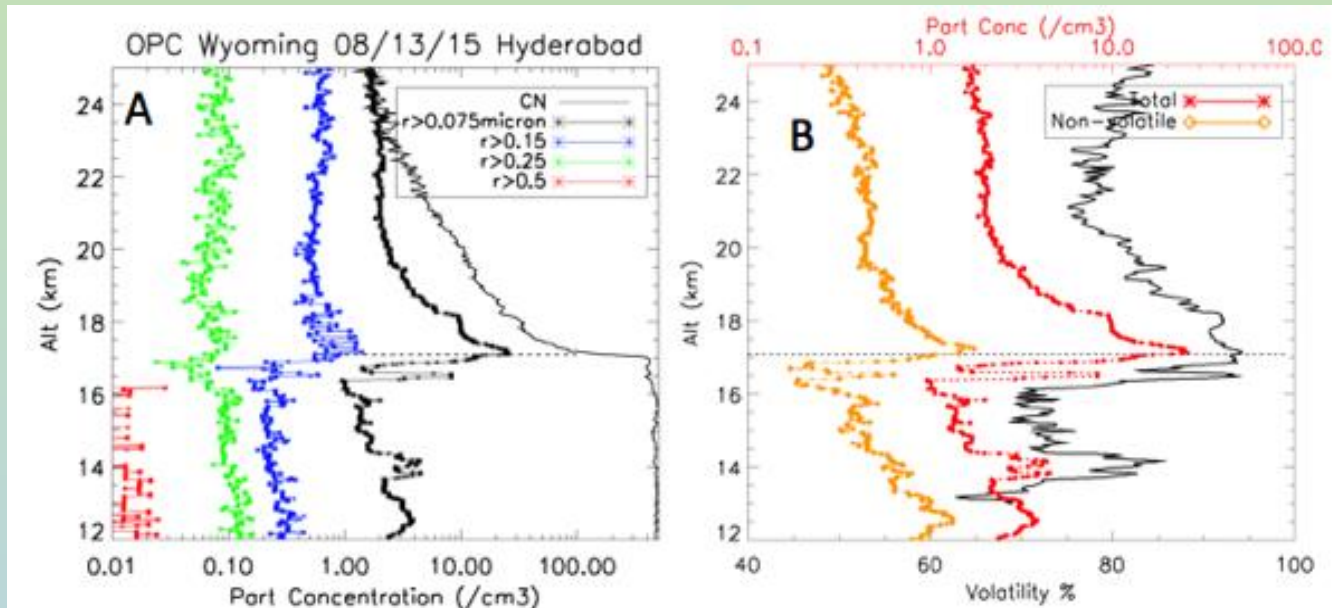


Vernier et al. (GRL, 2011)



Park et al., ACP, 2008

Aerosol counting measurements using balloons within the ATAL: BATAL campaigns (2/4)

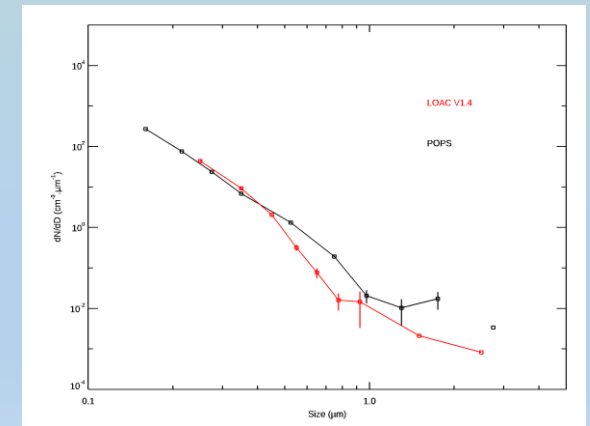
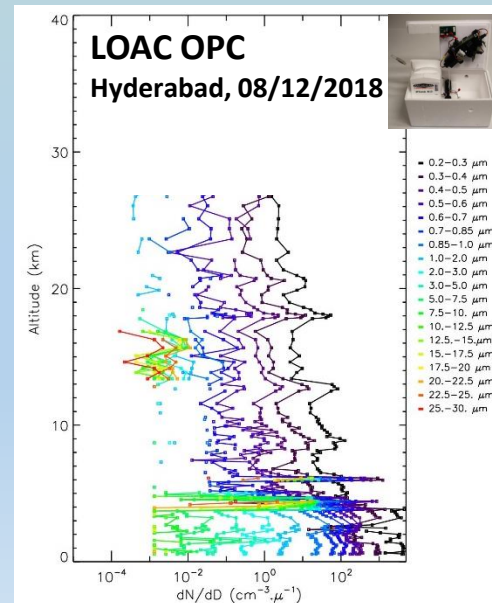
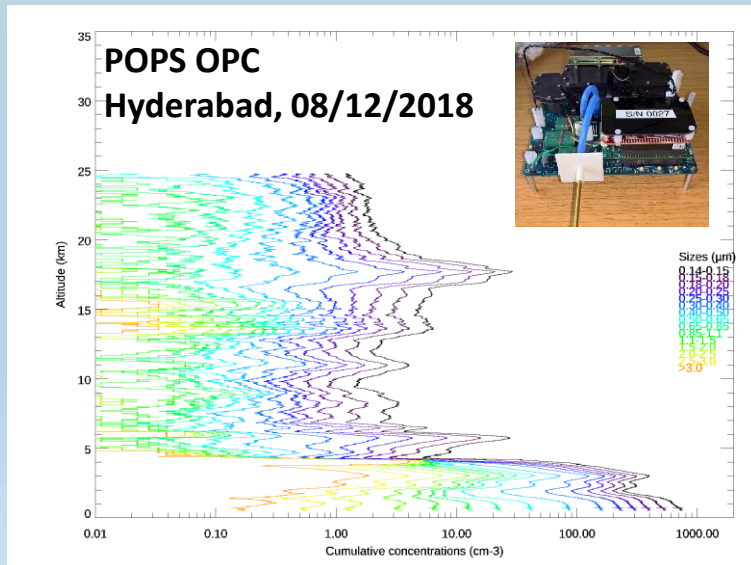


Vernier et al., BAMS, 2017



Campaigns coordinated by NASA, ISRO (India), NARL (India)

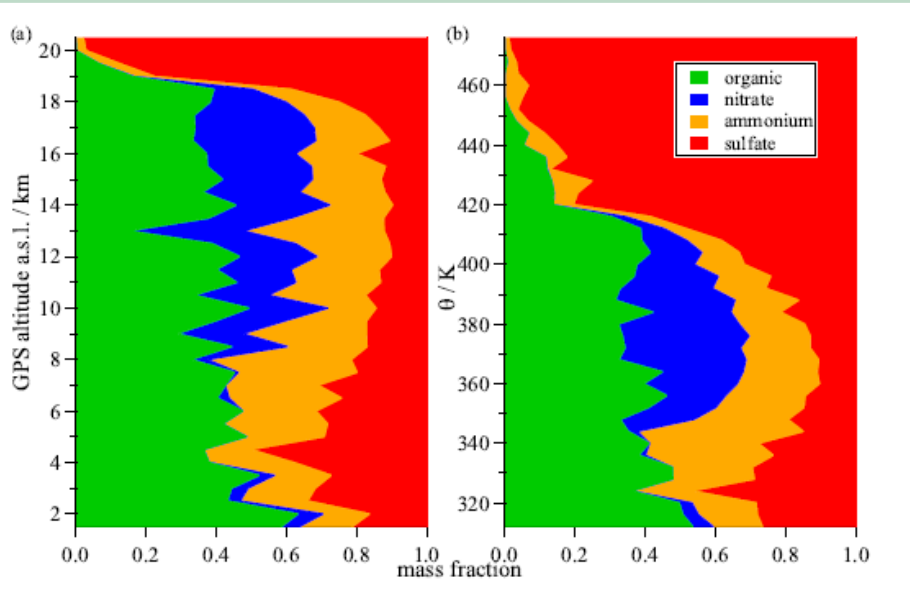
- Sharp increase of aerosol concentration near 16-18 km
- 90% of volatile aerosol
- small particles



-> ATAL made of very small, volatile aerosol and thus newly formed particles ?

Aerosol composition within the ATAL (3/4)

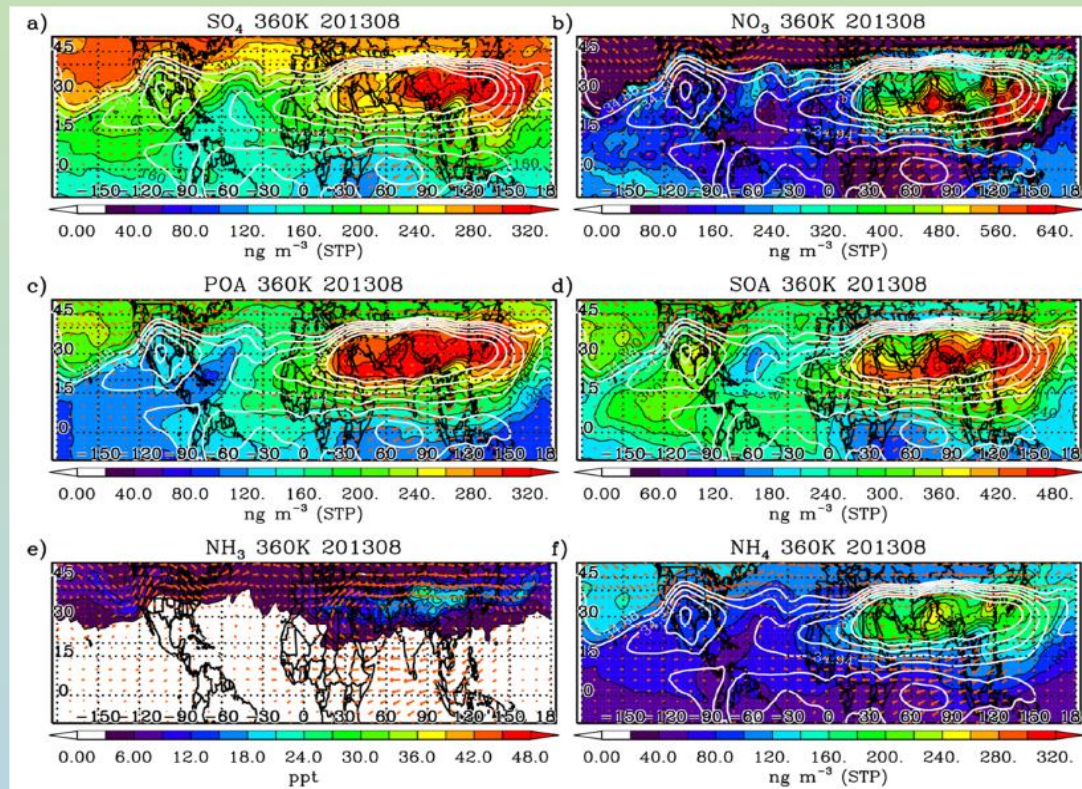
StratoClim EU campaign, summer 2017
 Mass spectrometry observations, M-55 Geophysica
 Nepal, India, Bangladesh, Bay of Bengal



Appel et al., ACPD, 2022

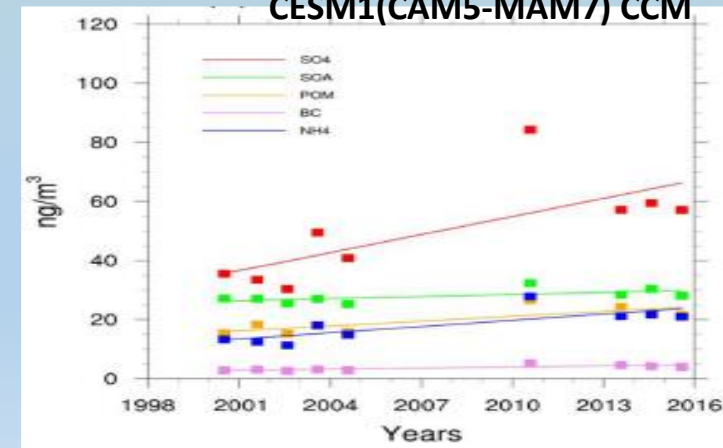
-> ATAL particles mainly resulting from the conversion of inorganic and organic gas precursors rather than from the uplift of primary particles from below.

GEOS-Chem CTM, summer 2013



Fairlie et al., JGR, 2019

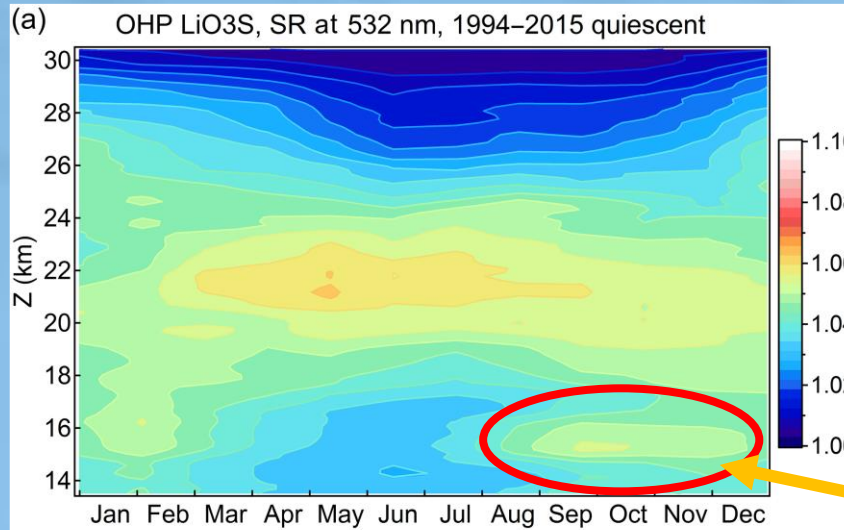
CESM1(CAM5-MAM7) CCM



Bossolasco et al., ACP, 2021

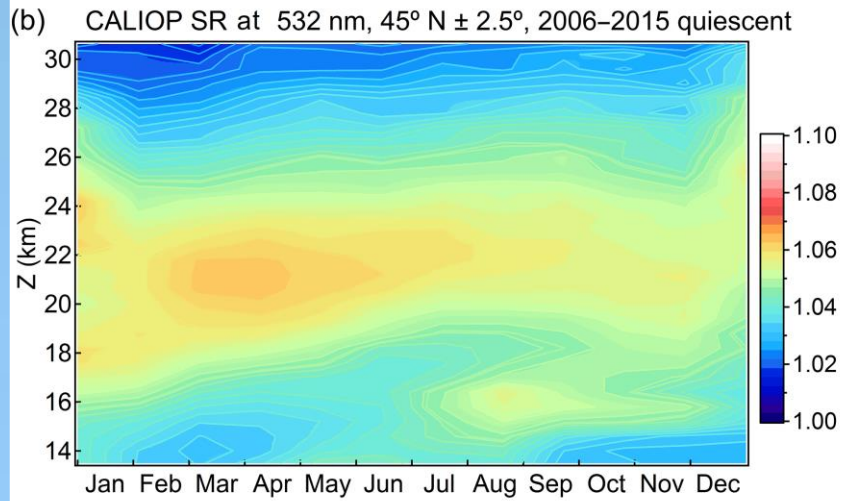
Signature at mid-latitudes (4/4)

Ground-based lidar
Observatory of Haute
Provence (44°N; 6°E),
France



SR = Scattering Ratio

CALIOP/Calipso
Zonal mean 45°N



*Signature of ATAL
aerosols above
France after the
breakdown of the
monsoon anticyclone*

VOLCANIC PLUMES

Interannual variability of the stratospheric aerosol content

Tropics 20°S-20°N

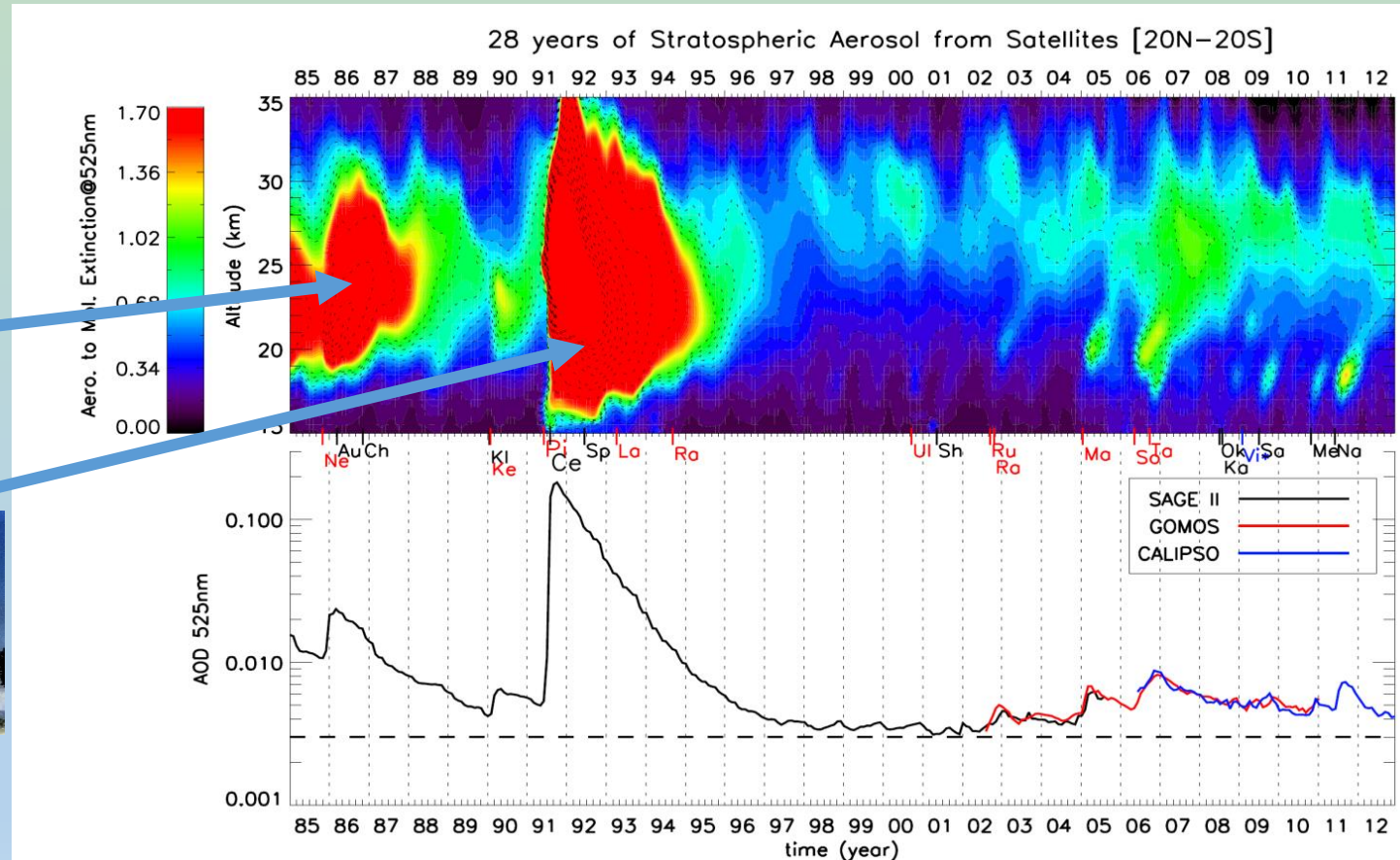
SAGEII UV-Visible spectrometer
CALIOP/Calipso space-borne lidar

Extinction ratio
(= Aerosol Extinction /
Molecular Extinction)

The 2 last major
eruptions:

El Chichon

Pinatubo

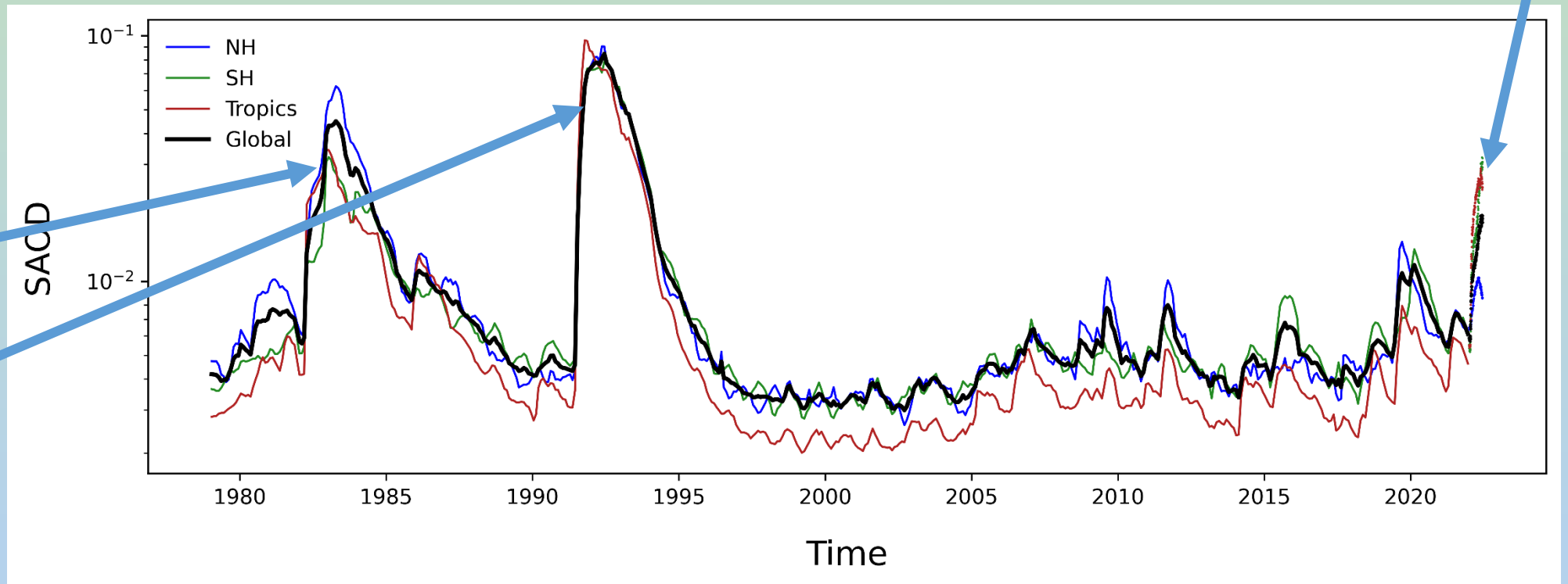


Courtesy from
Jean-Paul Vernier

Interannual variability of the stratospheric aerosol content

Merging of satellite data: GloSSAC v2.2 + OMPS-LP

Hunga Tonga



The 2 last major eruptions:

El Chichon

Pinatubo



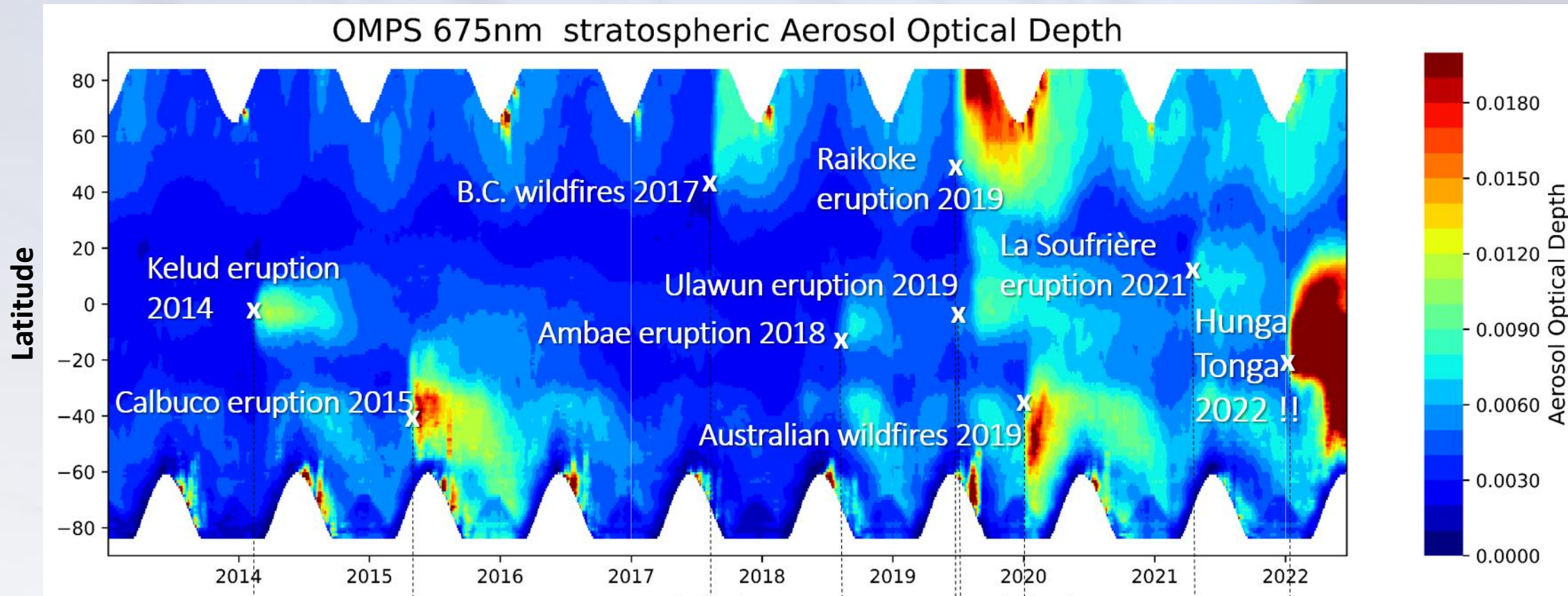
Courtesy from Pasquale Sellitto, Preliminary plot

Sellitto et al., under revision, Nature Comm., 2022, https://assets.researchsquare.com/files/rs-1562573/v1_covered.pdf?c=1650312853

doi: <https://doi.org/10.21203/rs.3.rs-1562573/v1>

Legras et al., ACPD, 2022, <https://doi.org/10.5194/egusphere-2022-517>

Over the more recent years

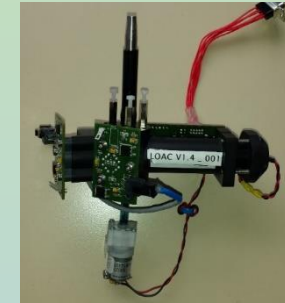


In situ observations of volcanic plumes (HEMERA support)

HEMERA frame and support for:

- instrumental development (e.g. LOAC V1.2 to V1.5 for improved performances: reduced stray light, powerful laser source)
- rare regular comparisons of in situ observations of stratospheric aerosols using small (latex) balloons: LOAC vs POPS + commercial instruments from NASA-LARC (J.-P. Vernier)
- Two flights in Kiruna in February and December 2020

Together with support from **University of Orleans** (VOLTAIRE project), **CNES**, **ANR** (French national Research Agency)

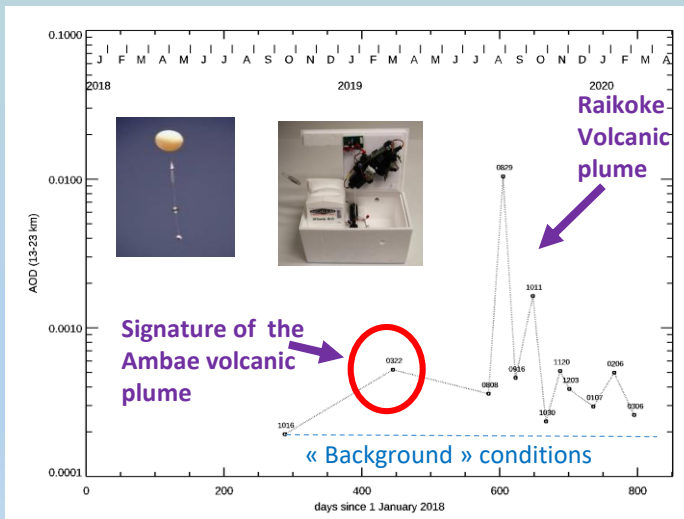


LOAC (LPC2E & MeteoModem company, France) 1 kg payload

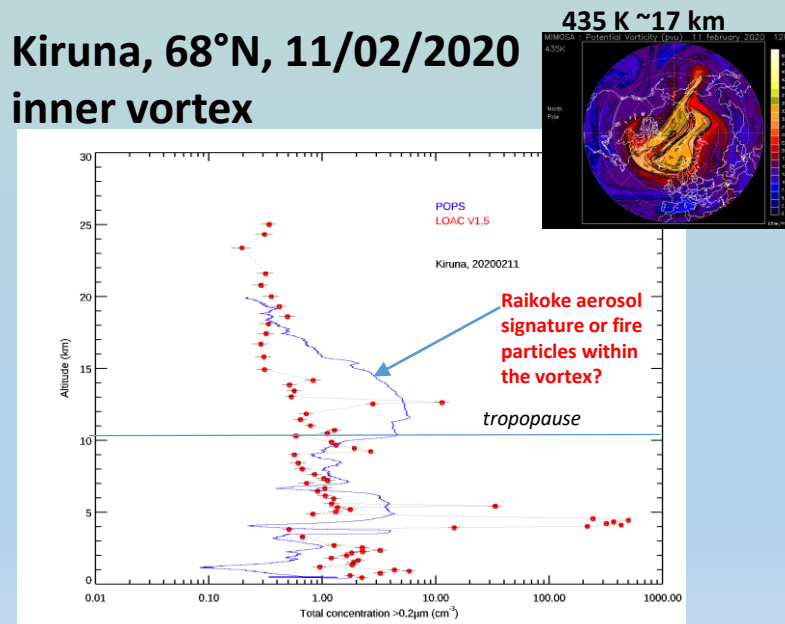


POPS (Handix Scientific company, USA) 1 kg payload

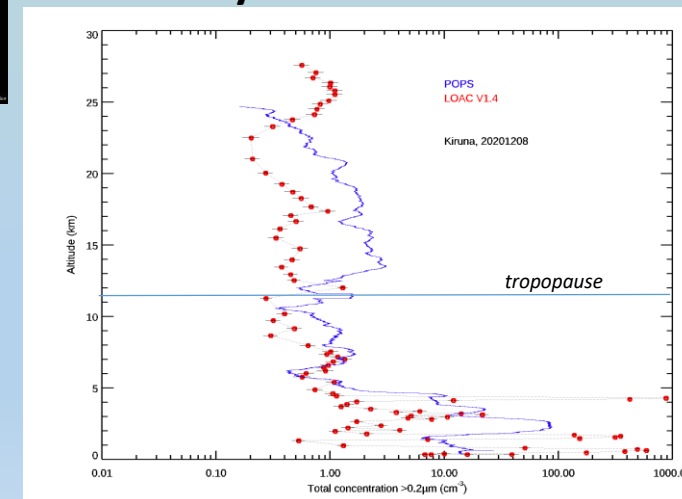
Mid-latitudes



Kiruna, 68°N, 11/02/2020 inner vortex



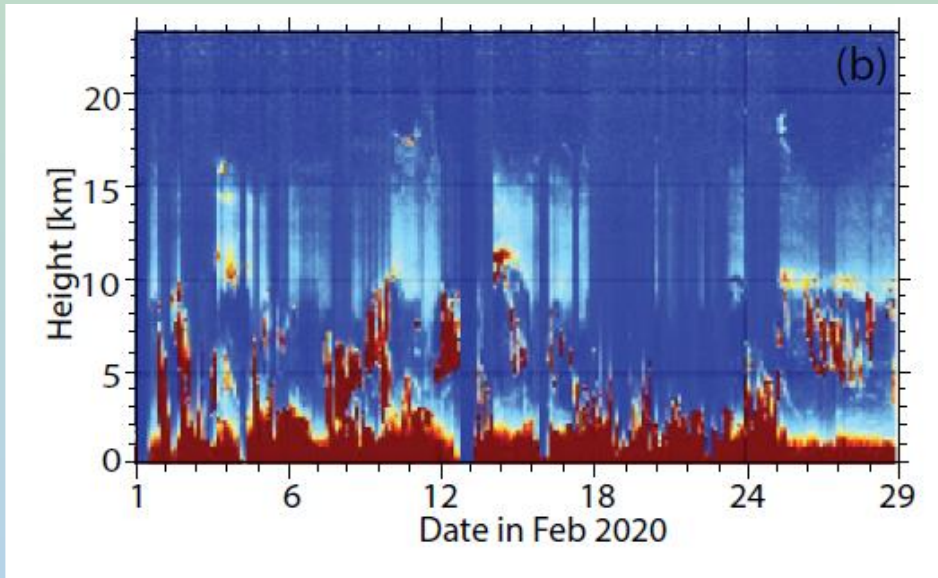
Kiruna, 68°N, 08/12/2020 early vortex



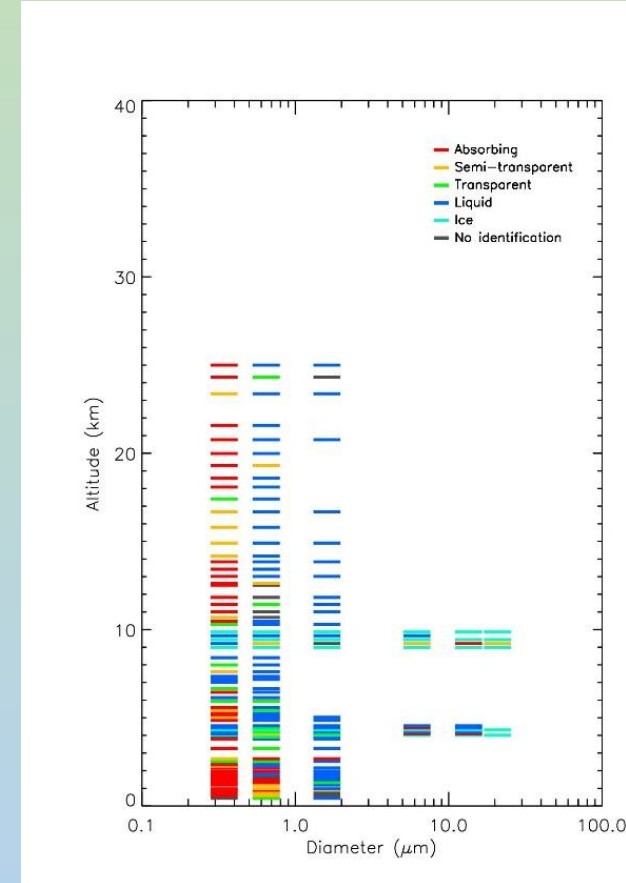
LOAC OPC flights from CNES base (Asa, 44°N) and MeteoModem company site (48°N), France

-> Impact of fire smoke on ozone depletion in winter 2019-2020 proposed by Ohneiser et al., ACP, 2021; Ansmann et al., ACPD, 2022)

Winter
2020



Lidar observations (of smoke haze?) over the Polarstern (85-88.5°N)
Ansmann, ACPD, 2022



LOAC v1.5 OPC
Kiruna, 68°N, 11/02/2020
inner vortex

Need for model
simulations including
volcanic/smoke plume

Community Aerosol Model for Atmospheres (CARMA) Coupled with the CESM1(WACCM) Chemistry-Climate Model

Simulations:

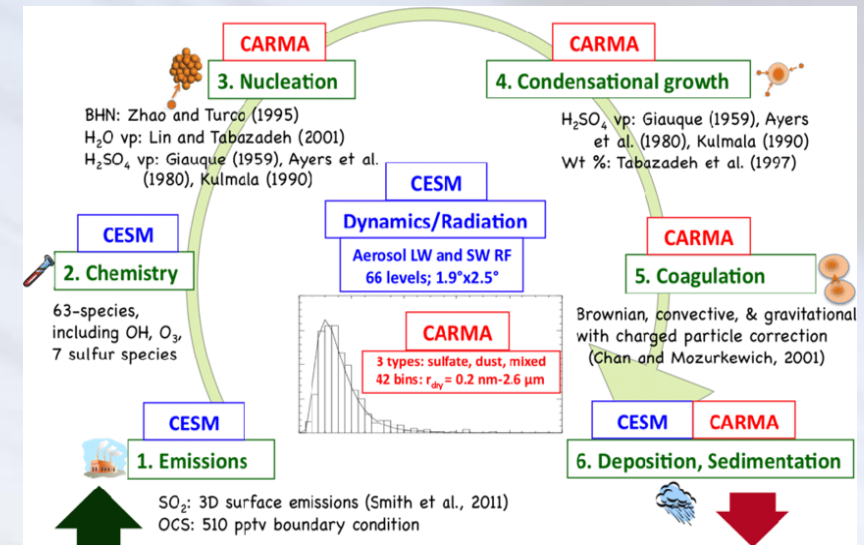
CESM1(WACCM) 1.9°x2.5° lon/lat, 88 vertical levels

Chemistry : waccm_mozart_sulfur

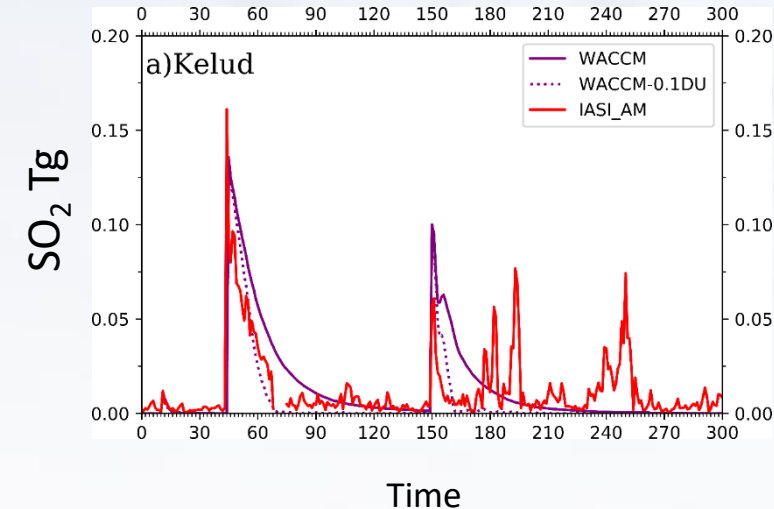
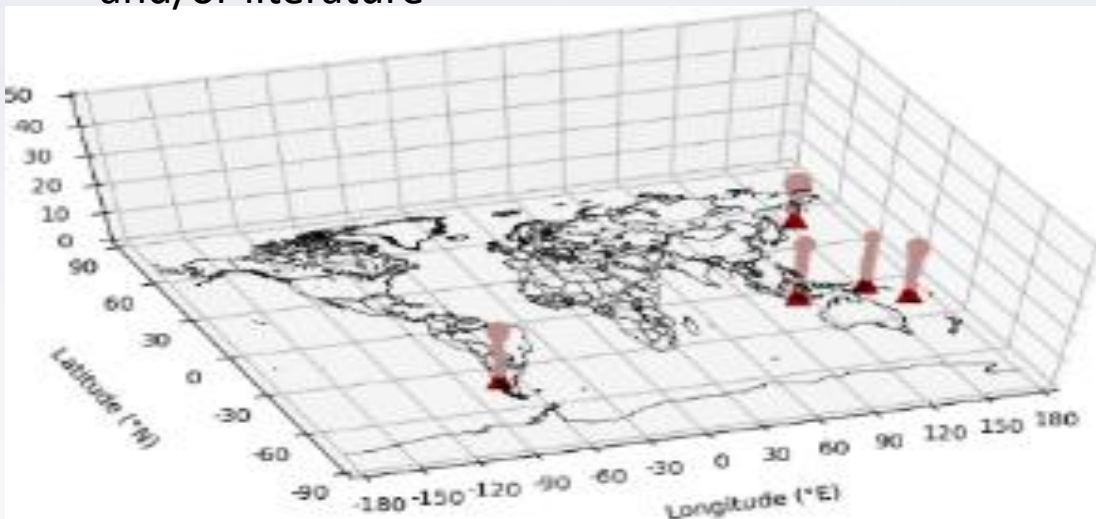
CARMA microphysical module : 30 aerosol size bins

Dynamics: Nudging towards MERRA2 reanalysis

Volcanic emissions of SO₂ (injection timing, altitude, amount of sulfur): model driven by information from satellite data and/or literature



English et al., ACP, 2011

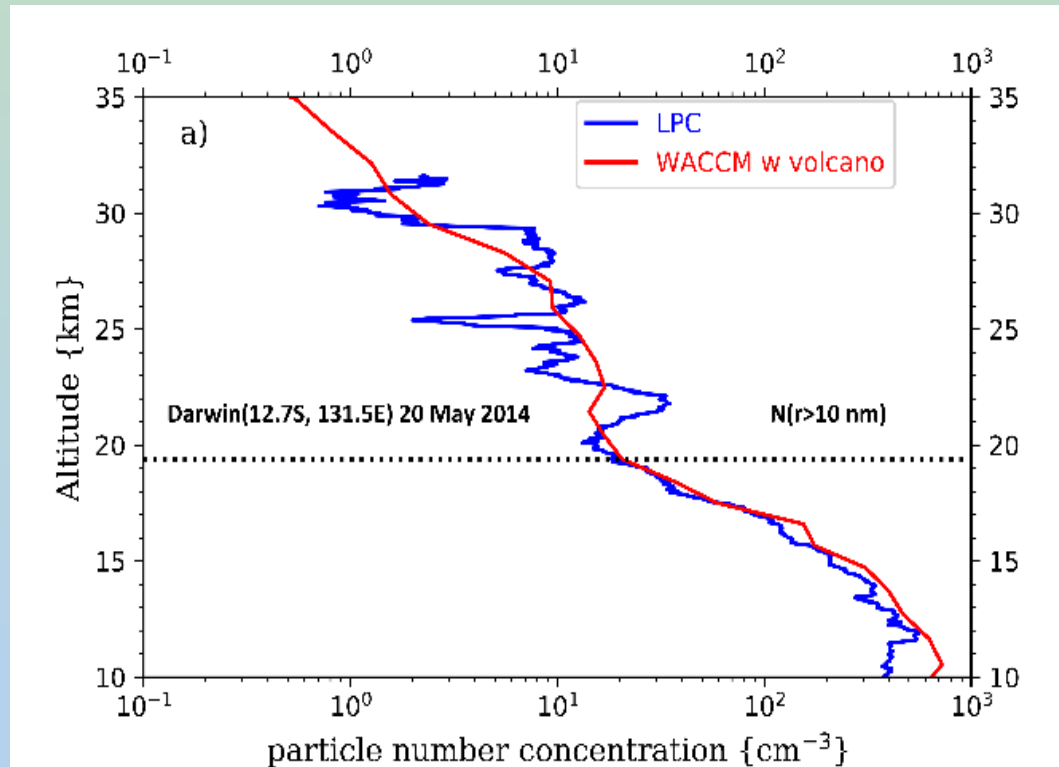


Tidiga et al.,
Atmosphere,
2022

Simulating the condensation nuclei profiles

LPC Optical particle counter (Univ. Wyoming; PI: Terry Deshler) vs WACCM-CARMA model

Darwin, Australia
KIAsh campaign
May 2014

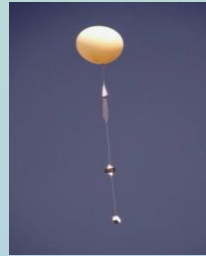
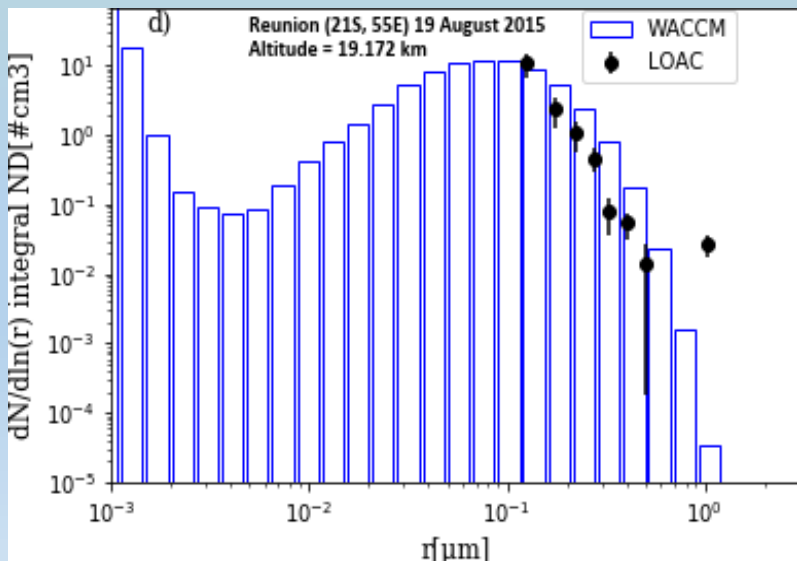
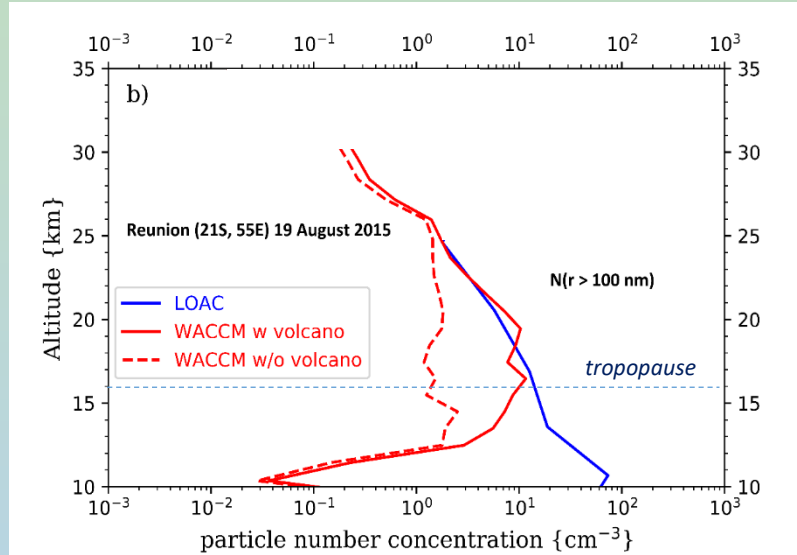


Tidiga et al., Atmosphere, 2022

Simulated vs observed size distributions of volcanic sulfuric acid particles

LOAC (v1.2) OPC vs WACCM-CARMA model
La Réunion island.

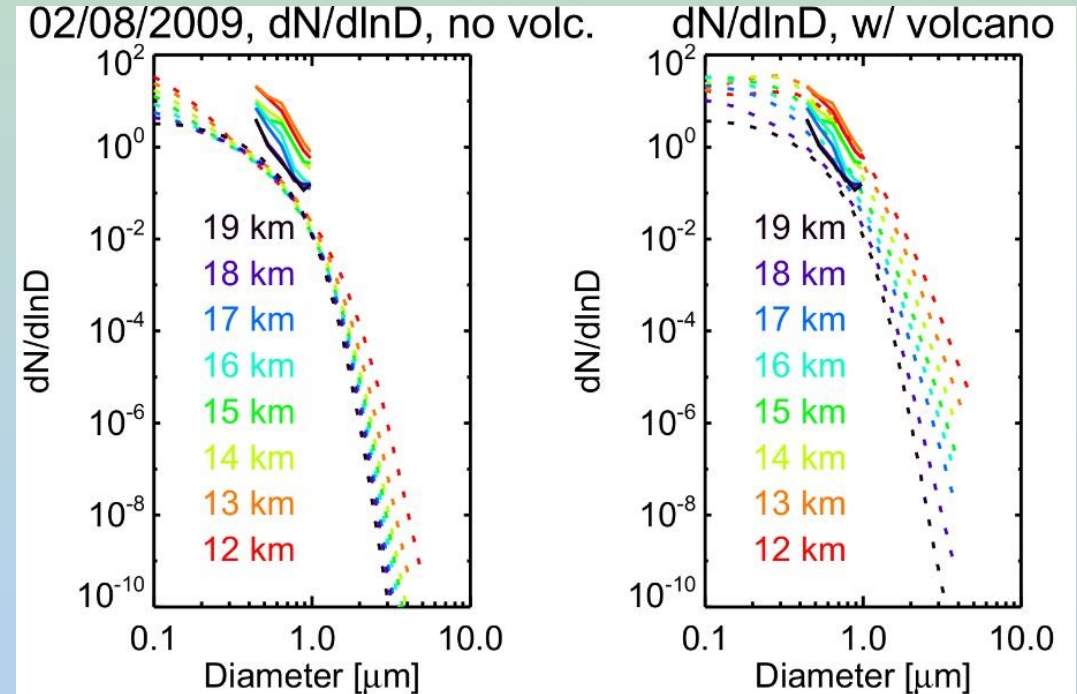
Calbuco volcanic plume (2015)



Adapted from
Tidiga et al., Atmosphere, 2022
and *Zhu et al., JGR, 2018*



STAC OPC vs WACCM-CARMA model, Kiruna, StraPoLEté project, Northern Sweden. **Sarychev volcanic plume (2009)**



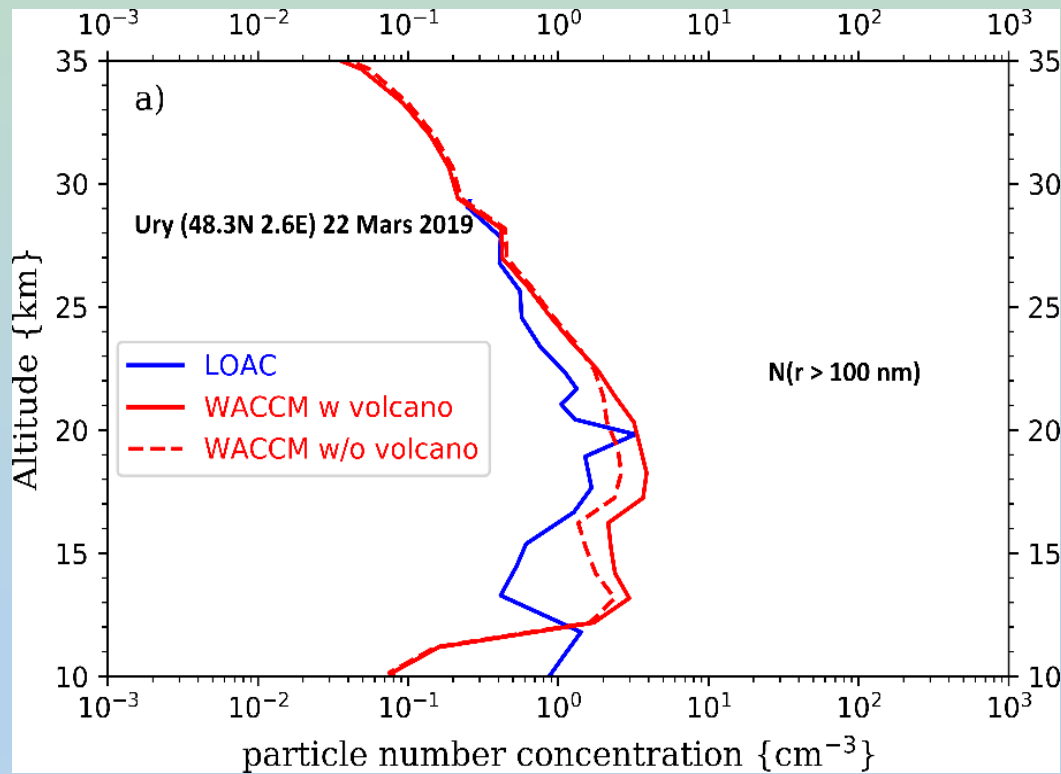
Dotted Lines= WACCM-CARMA model

Full lines = Optical Particle Counter observations

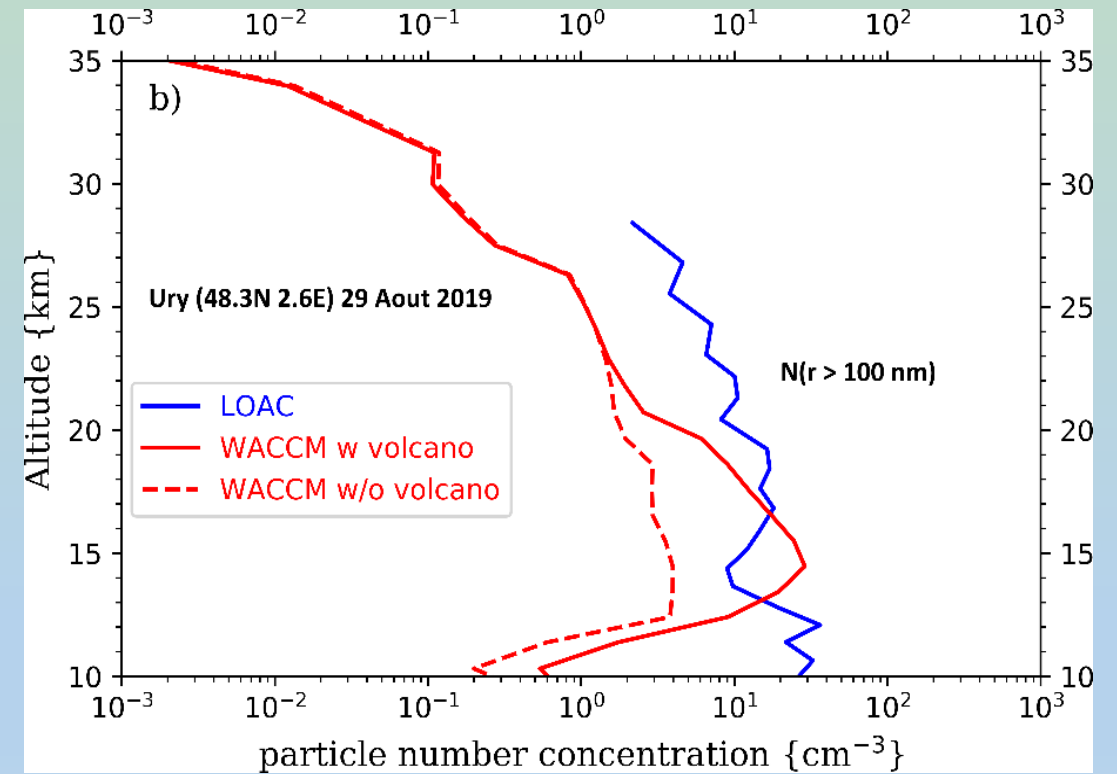
Lurton et al., ACP, 2018

Simulated vs observed size distributions of volcanic particles

« Aged » Ambae plume
(eruption 07/2018)
Decaying in 2019



Early Raikoke plume
(eruption 06/2019)



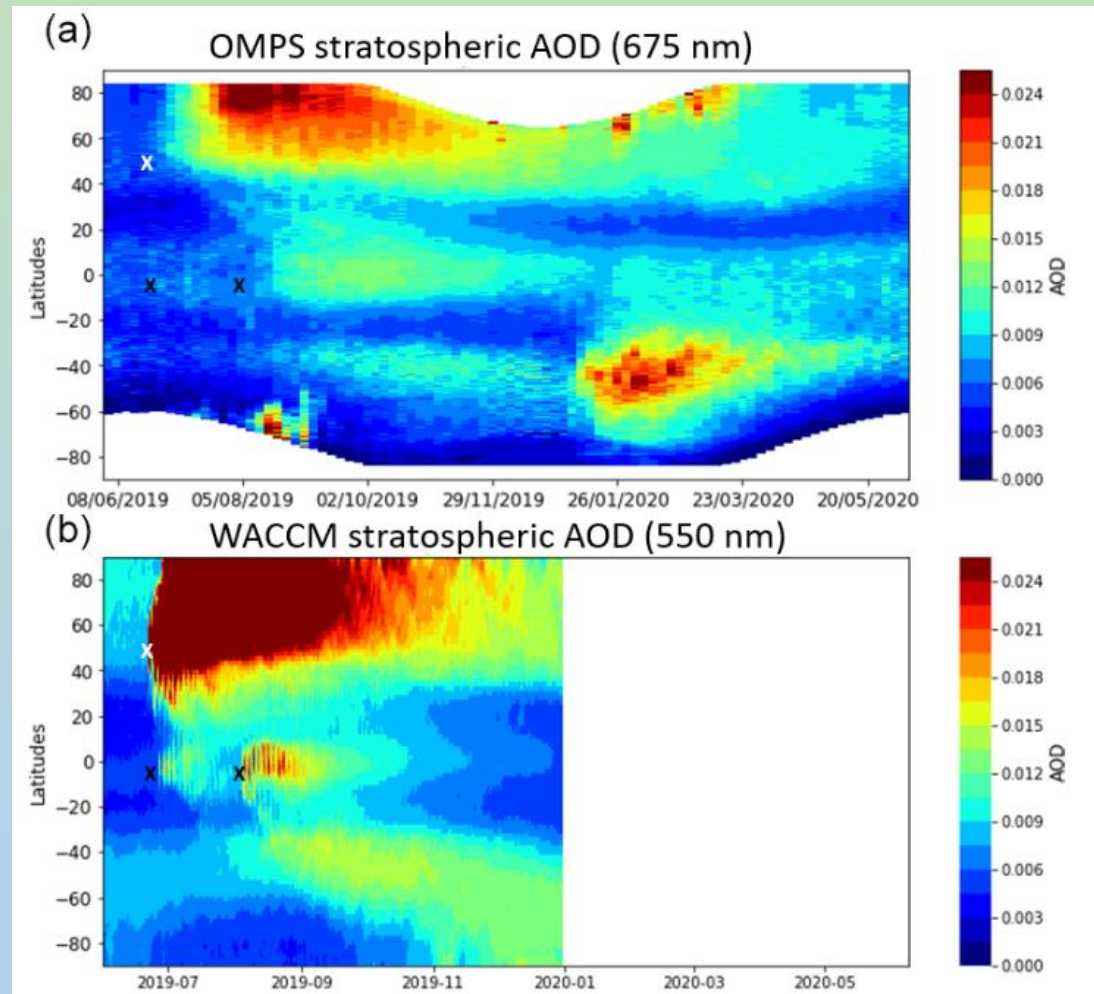
Simulating the transport of the Raikoke plume

Aerosol Optical Depth

OMPS

vs

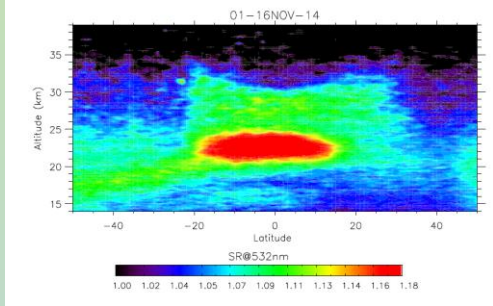
WACCM-CARMA model



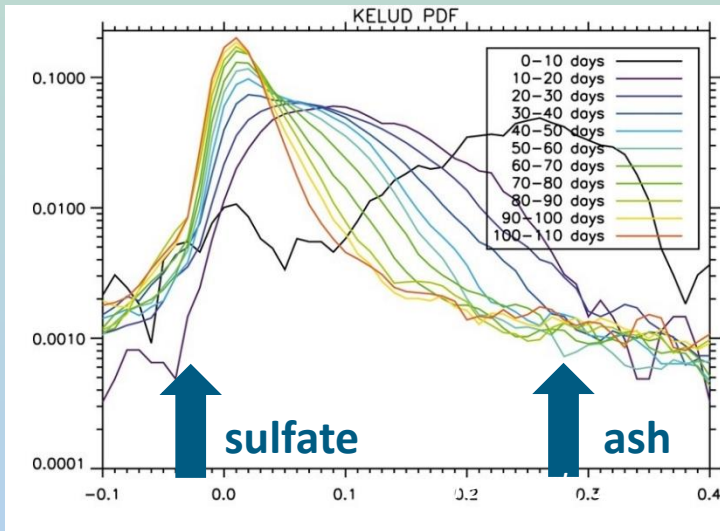
Kloss et al., ACP, 2021

Injection parameters (SO_2 burden, altitude range, injection timing) not always derived robustly from satellite observations

Co-injection of ash



Ash detection by CALIOP/Calipso Kelud volcano case (Feb. 2014)



Depolarization ratio

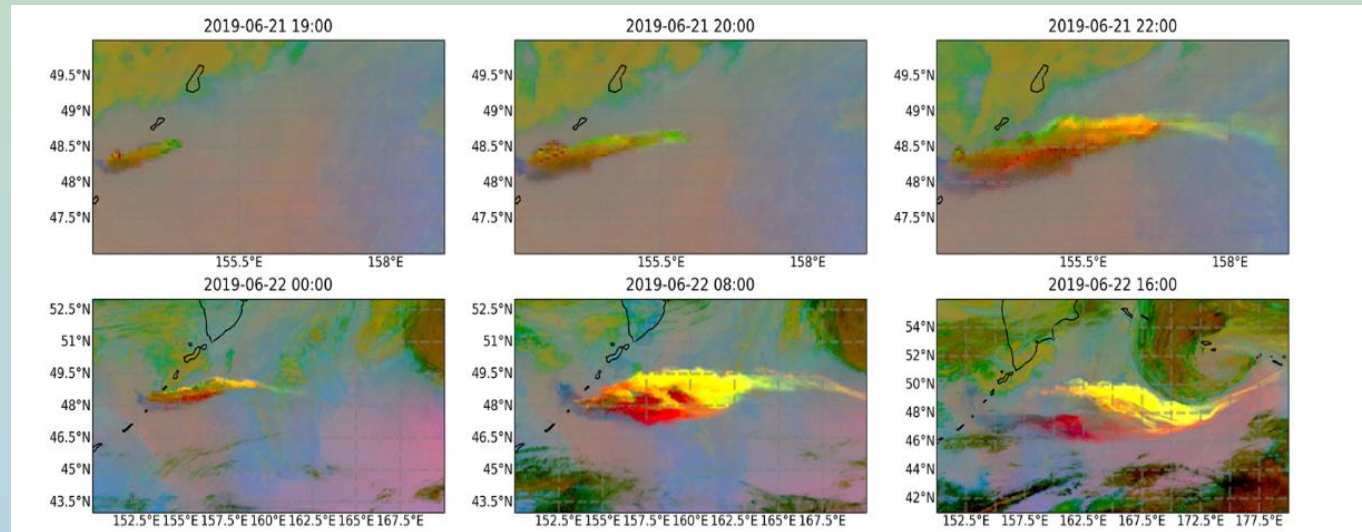
Vernier et al., JGR, 2016

Raikoke volcano, Kuril Islands (north of Japan)

Himawari RGB images
June 2019



Source - © 2019 NASA / Earth Observatory



Red: ash Green: sulfur Yellow: mixture

Kloss et al., ACP, 2021

Radiative effects of ash on the plume transport (see Muser et al., ACP, 2020)

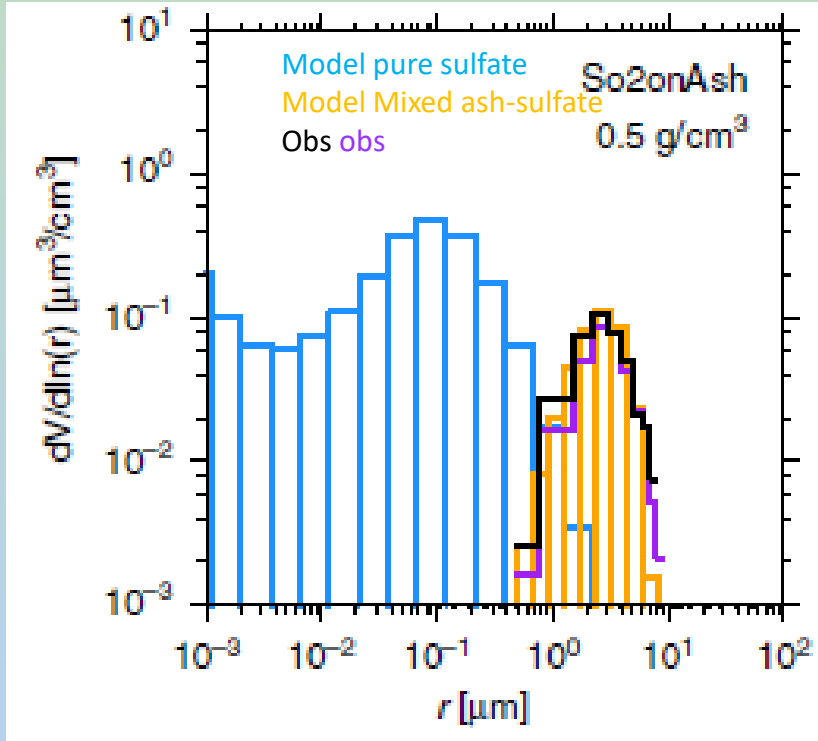
+ injection of water vapour?

+ interaction with smoke plumes from north America?

Chemical role of ash

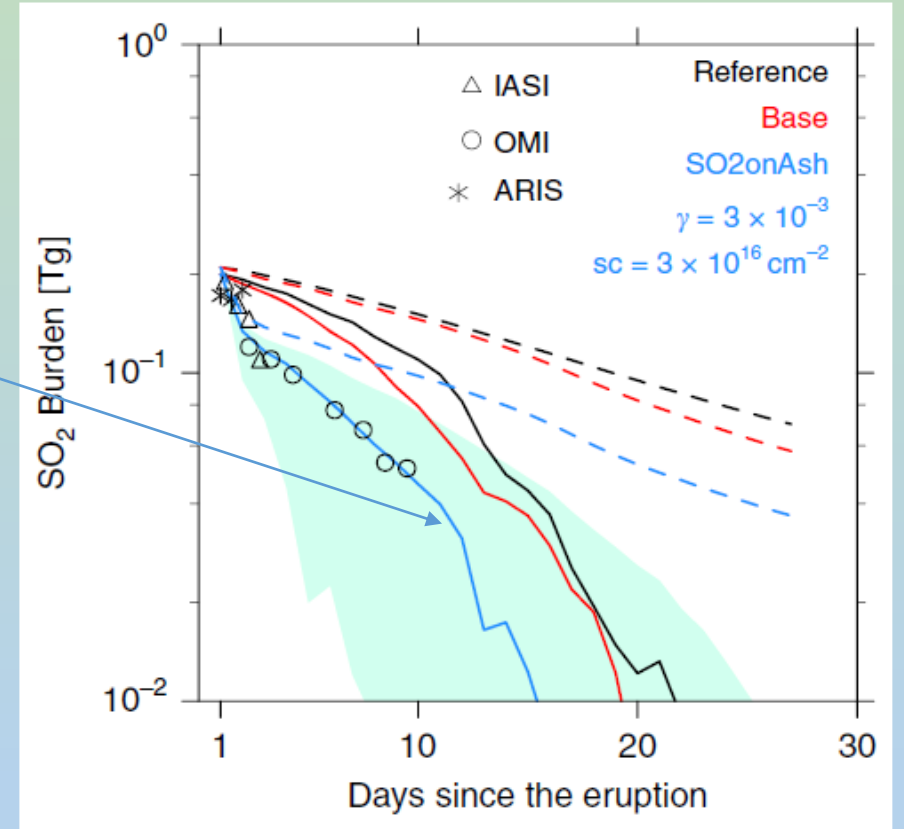
Kelud volcanic plume (2014)

Above Guam Island
16.5-18.5 km, March 2014



With SO₂ uptake on ash

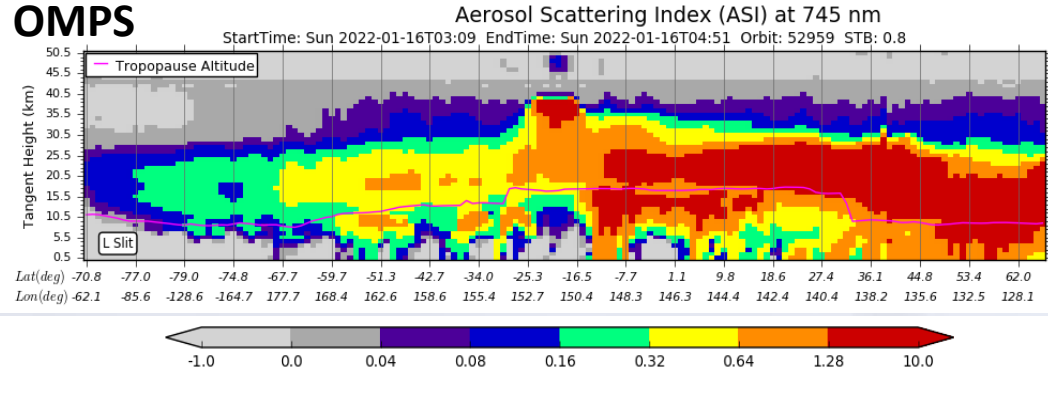
Zhu et al., Nature Comm., 2020



-> Ash tends to decrease SO₂ lifetime

Impact of Hunga Tonga eruption (January 2022)

La Réunion island



LOAC V1.5 flights from La Réunion island under alert

Support: HEMERA and VOLTAIRE project (Univ. of Orleans)

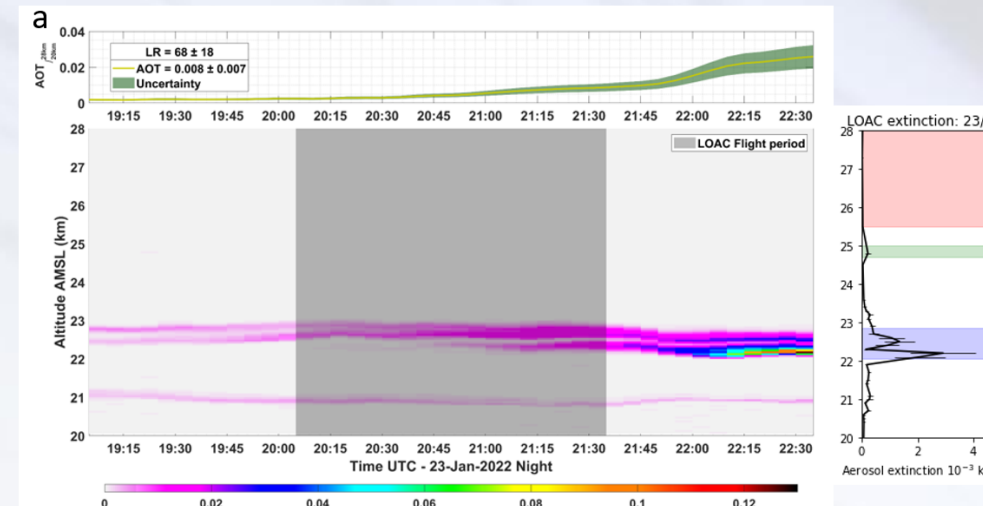
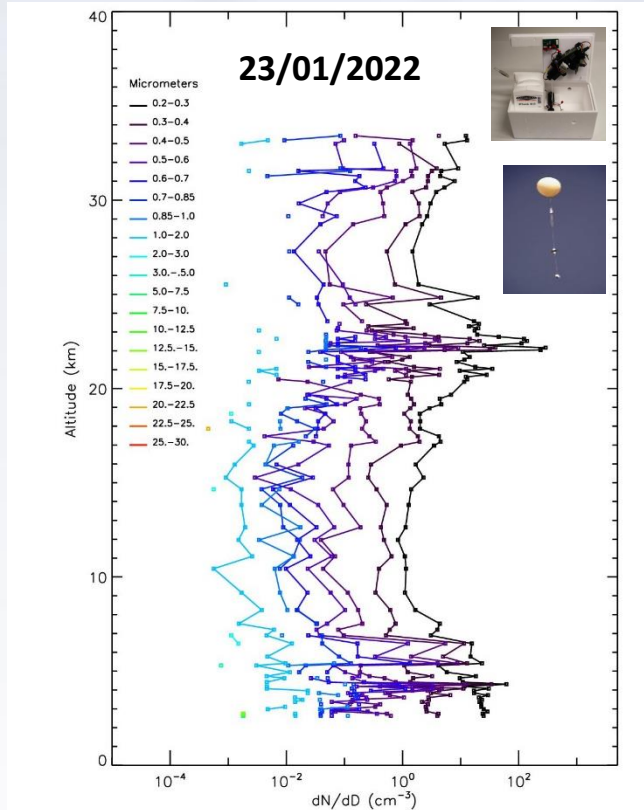
LOAC V1.5



<https://la1ere.francetvinfo.fr>



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Optically-absorbing features of the small particles

Kloss et al., under revision, 2022

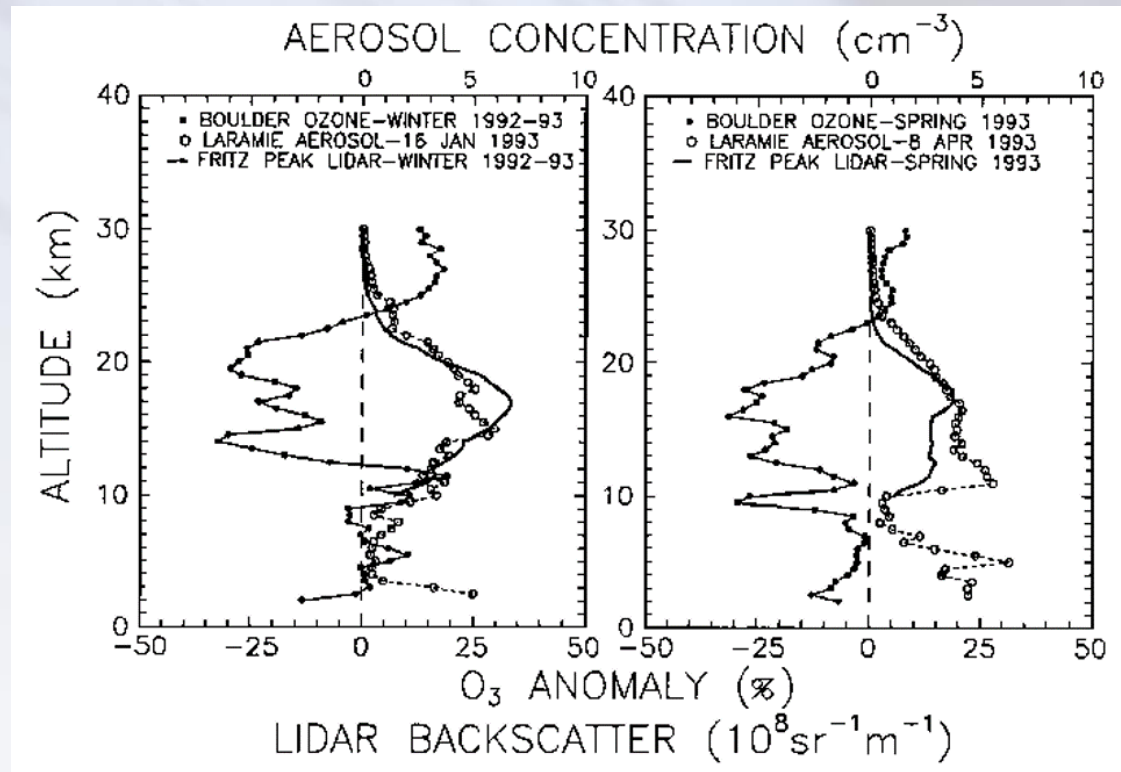
Eruption with unique characteristics:

- Injection **above 40 km** (signal at **50 km!**)
- Injection of **ash**, **halogens**, huge quantities of **water vapour** with very strong impact on sulfur chemistry, production and evolution of the aerosols, and stratospheric ozone chemistry

Ozone loss at mid-latitude after the Pinatubo 1991 eruption



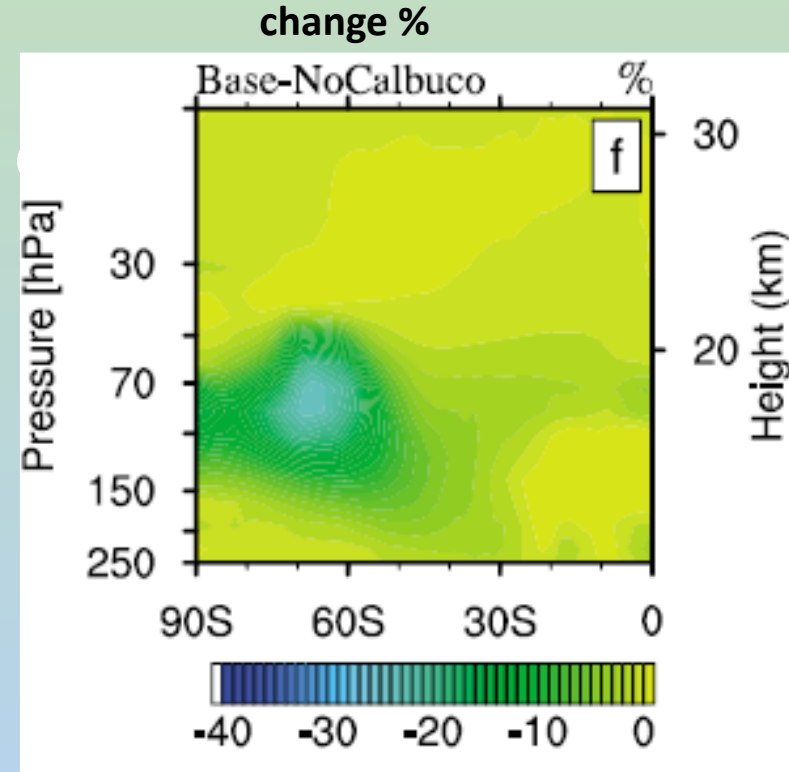
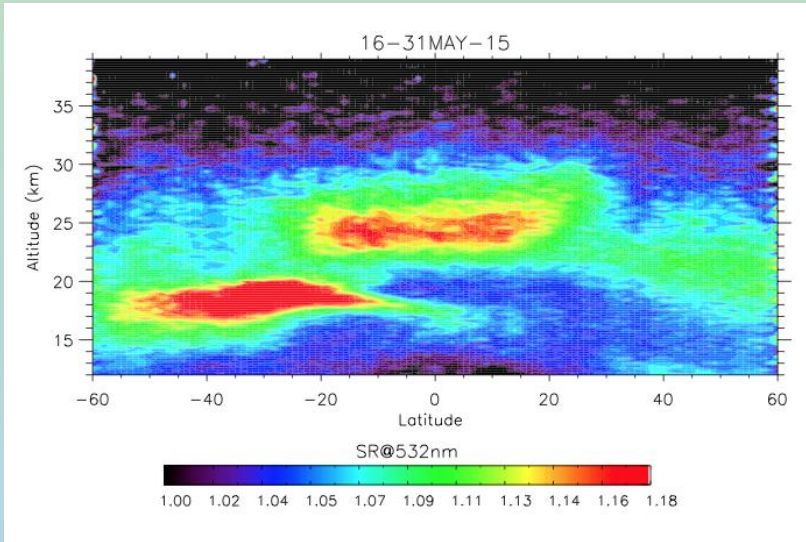
~20 Tg of SO₂ injected



Hoffmann et al. GRL, 1994

Volcanic aerosols from moderate eruptions: chemical impacts

Observations by CALIOP/Calipso
Scattering ratio (SR)



Impact of Calbuco volcano aerosols

CESM1(WACCM)-CARMA model

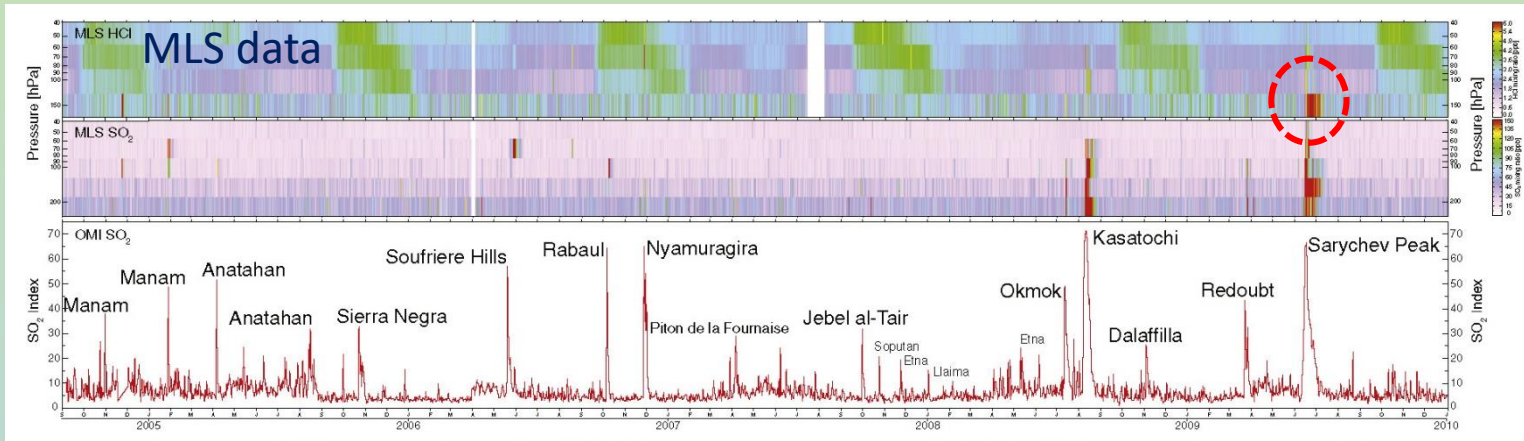
Antarctic polar vortex

15-30 September 2015

Zhu et al., JGR, 2018

→ ~25% of further ozone depletion (zonal mean) due to volcanic aerosol presence in the Antarctic polar vortex

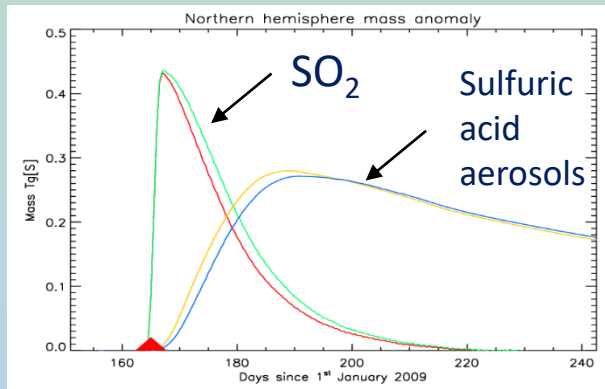
Co-injection of halogens



Northern hemisphere

HCl injection (27 Gg)

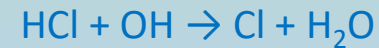
Carn et al., J. of Volc. And Geoth. Res., 2016



Sarychev eruption

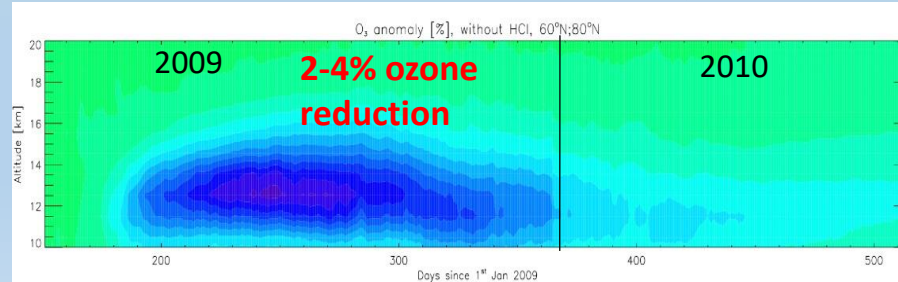
HCl injection → Slow down in the oxidation of SO₂ and sulfate aerosol formation: ~2-day delay

Combined effects of:

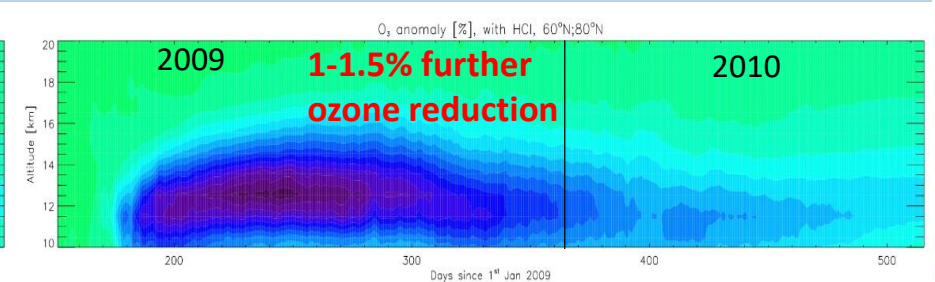


Lurton et al., ACP, 2018

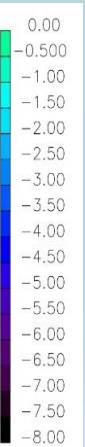
WACCM-CARMA CCM, no HCl injection



WACCM-CARMA CCM, HCl injection



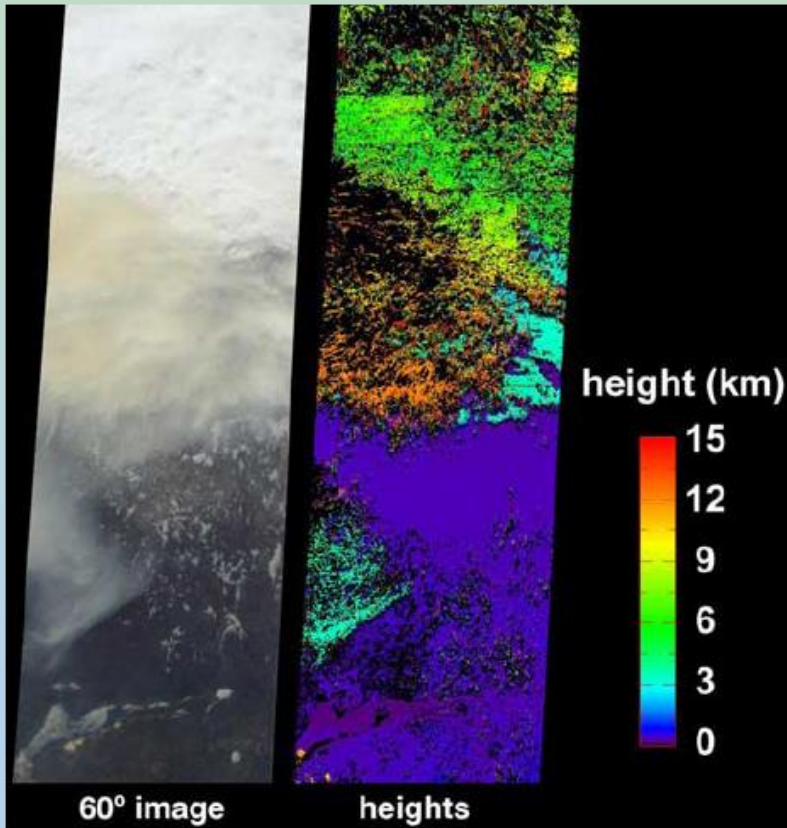
Linked to enhancement of HOCl in the Arctic polar vortex (+1 pptv as a result of cold temperatures)



FIRE SMOKE

Regular Pyrocumulonimbus (PyroCb) injections in the UTLS region

MISR images May 2001,
Alberta, Canada



Fromm et al., JGR, 2008

See *Fromm et al., BAMS, 2010* for a review

For example (strong events):

Norman Wells pyroCb, August **1998**, NW Canada

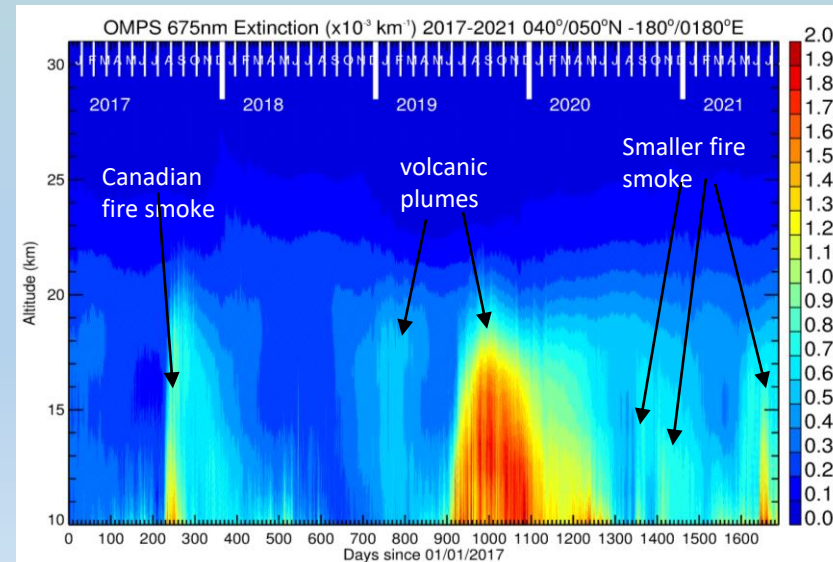
Chilsom fire, May **2001**, Alberta, Canada

PyroCbs of January **2003**, SE Australia

Black Saturday, February **2009**, Australia

Canadian fires, August **2017**

Australian bushfires, December **2019**-January **2020**

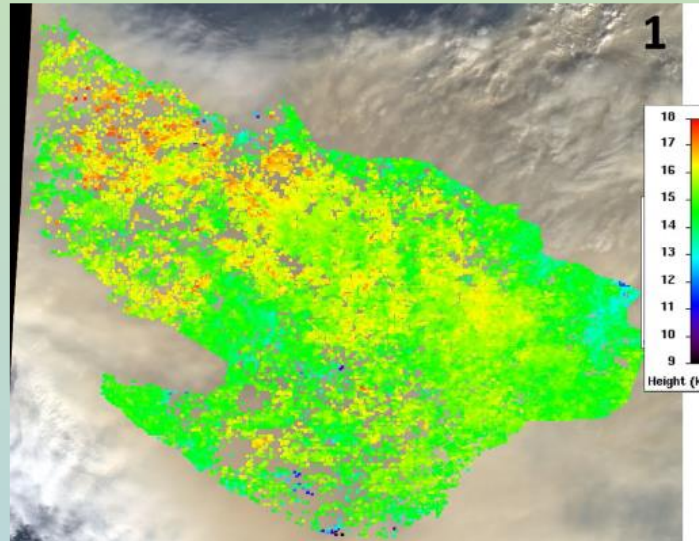


OMPS midlatitudes

**+ smaller
smoke injection
events**

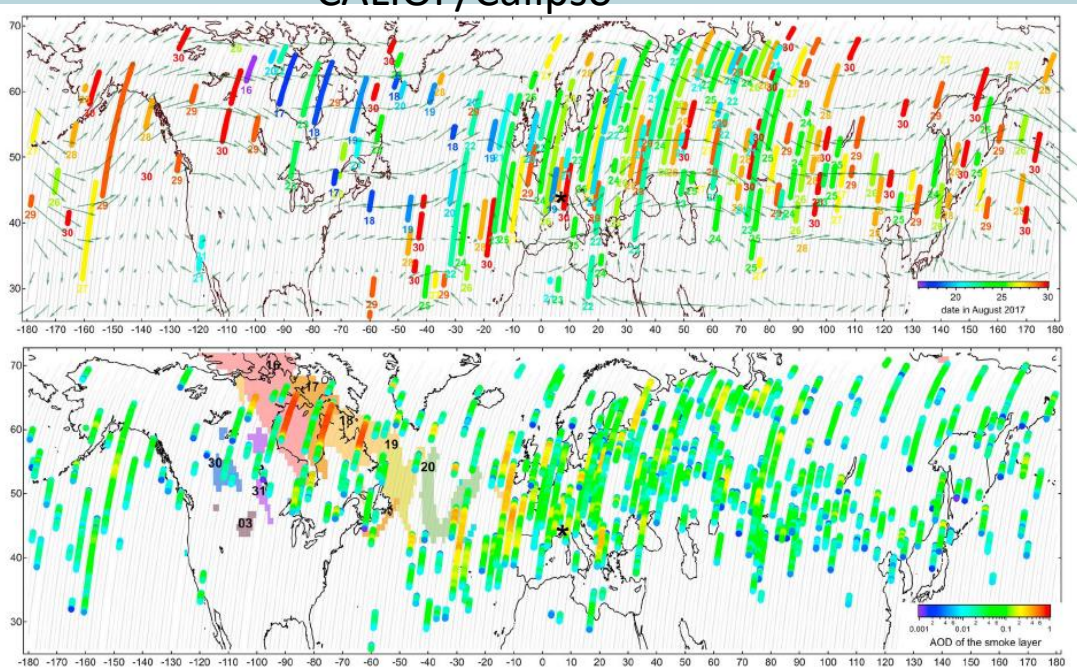
Plume transport: example of the Canadian fires, August 2017

MISR red-band data
Source region



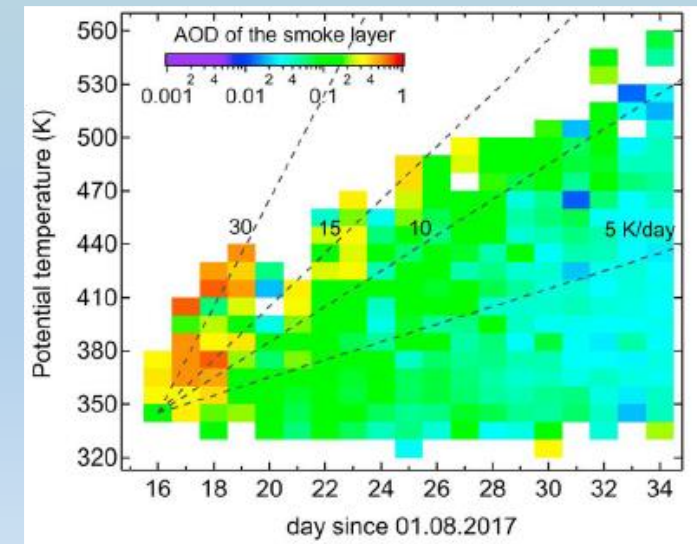
Fromm et al., JGR, 2021

CALIOP/Calipso



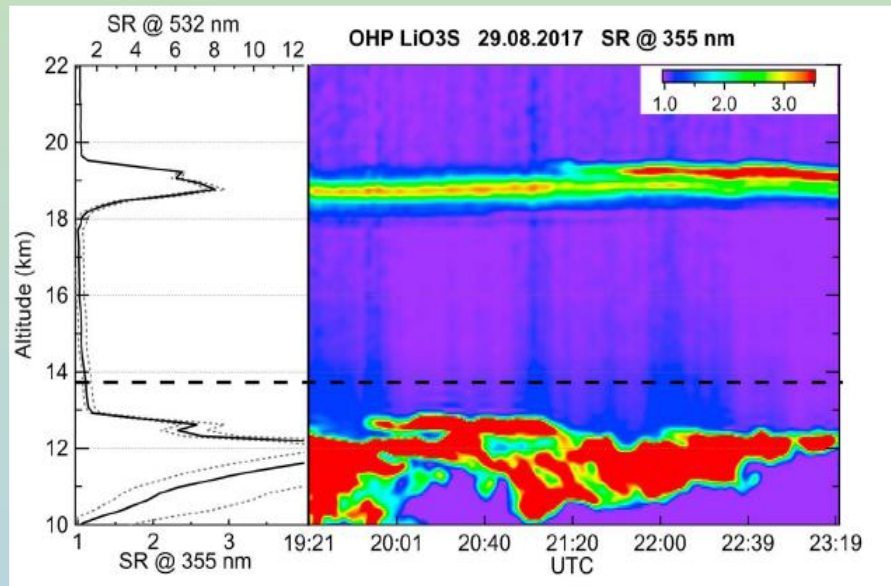
Khaykin et al., GRL, 2018

Plume rise consistent with localized heating due to the radiative absorption by the smoke particles

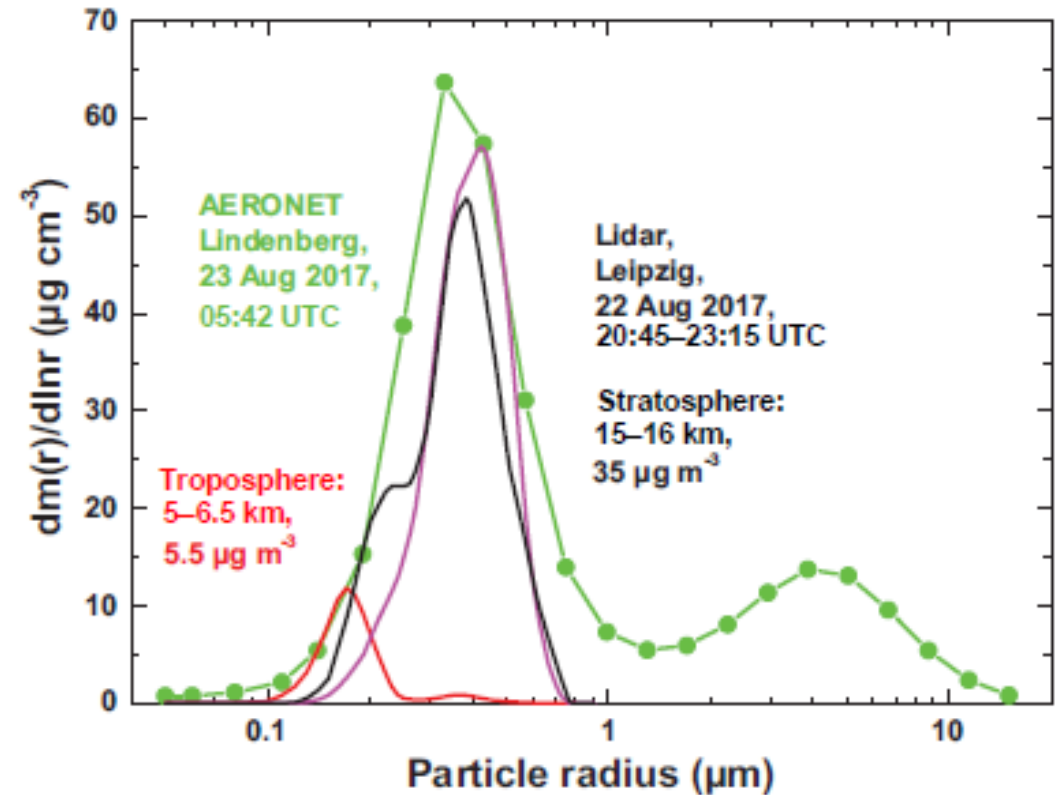


Information from ground-based observations

OHP lidar, France, 29/08/2017



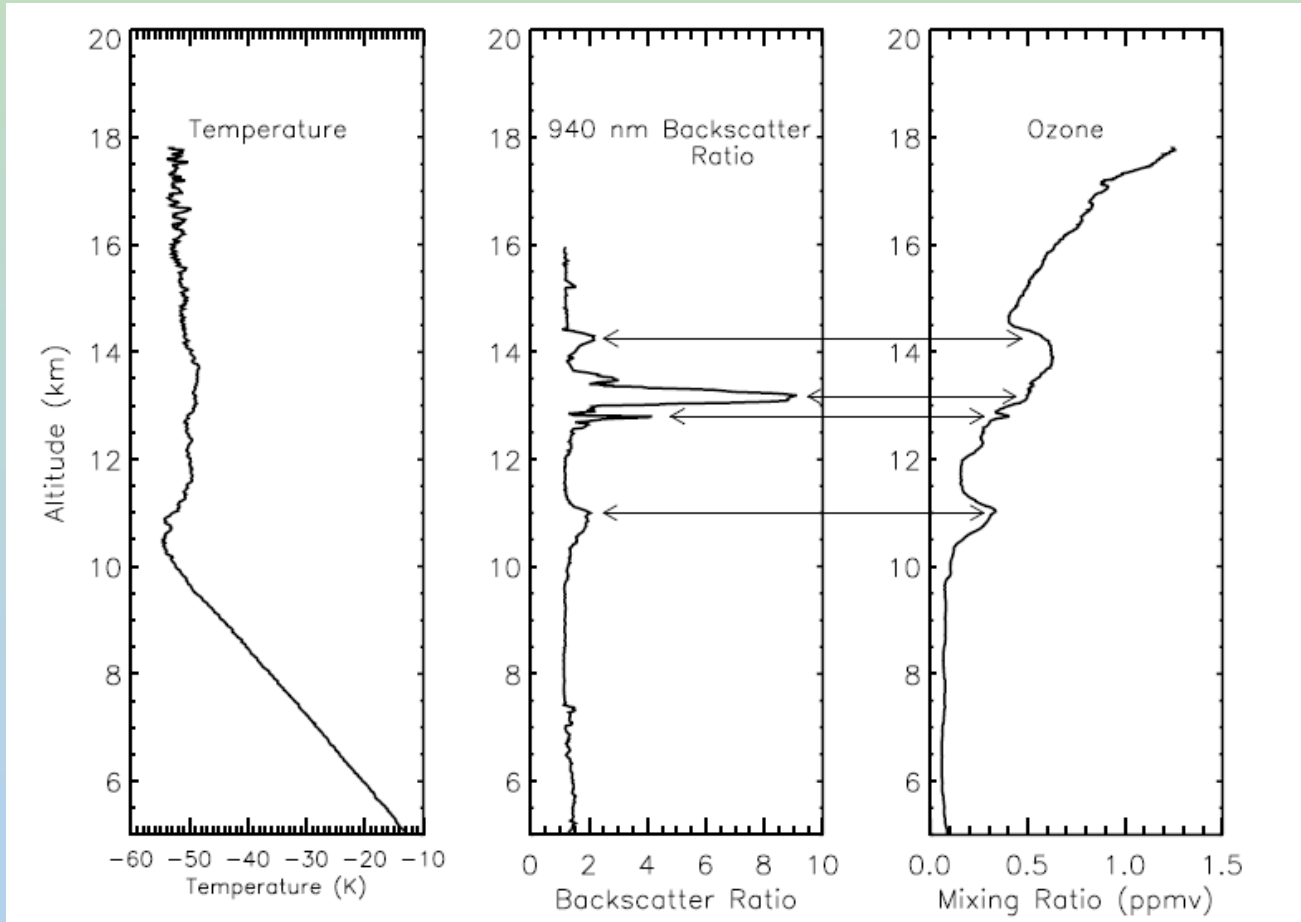
Khaykin et al., GRL, 2018



Haarig et al., ACP, 2018

- High depolarization ratios of 22% at 355 nm and 18% at 532 nm
- Pronounced accumulation mode (particle mass) centered at a particle radius of 0.35–0.40 μm .
- Effective radius : 0.32 μm (stratosphere), 0.17 μm (troposphere)
- Mass concentrations of $\sim 5.5 \mu\text{g.m}^{-3}$ (tropospheric layer) and $\sim 40 \mu\text{g.m}^{-3}$ (stratospheric layer) on 22/08/2017

Impact on ozone



Ozone production

depending on the levels of
NO_x, VOCs, HO_x

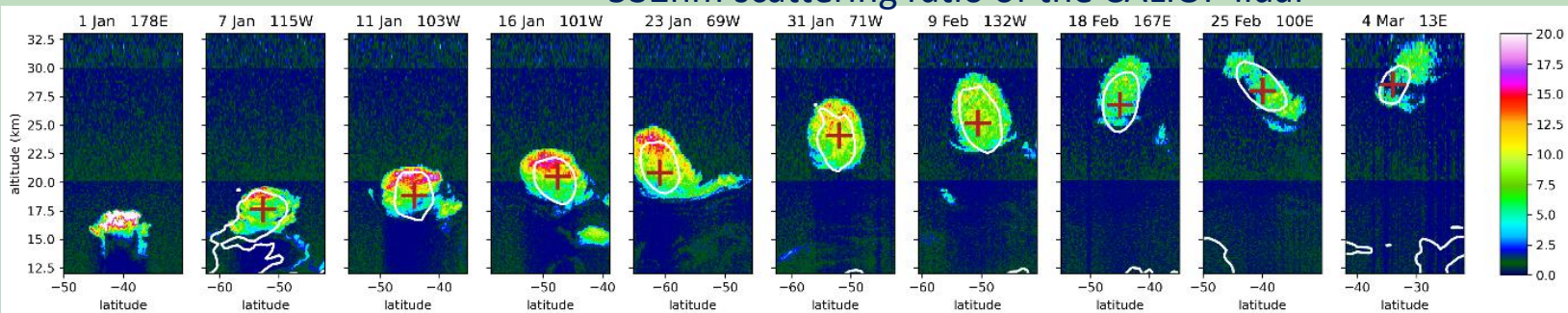
But what is the detailed
chemistry?

August 1998, Canada
(52°N; 107°W)

Self-maintained anticyclonic vortex of smoke particles

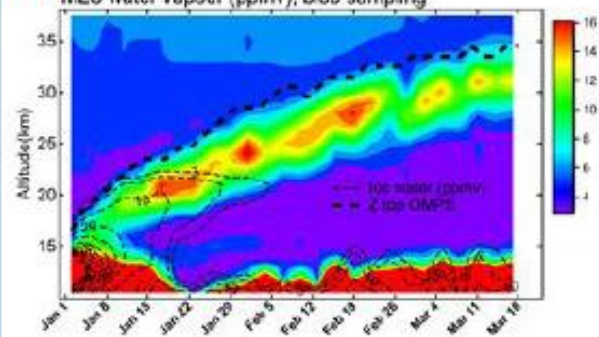
Australian fires in 2019-2020 but also observed for the Canadian fires in 2017 (see *Lestrelin et al., ACP, 2021*)

532nm scattering ratio of the CALIOP lidar

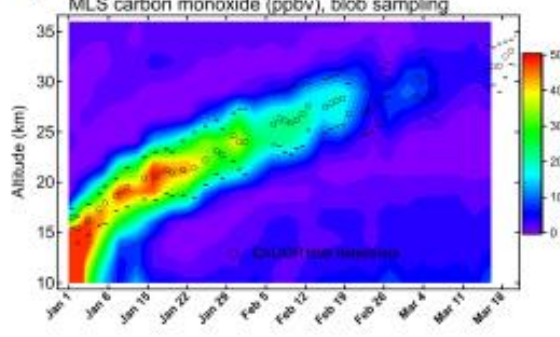


1,000 km diameter
Persistence over 13 weeks, 66,000 km travel
Confined bubble of smoke and moisture
+ shorter-lived companions

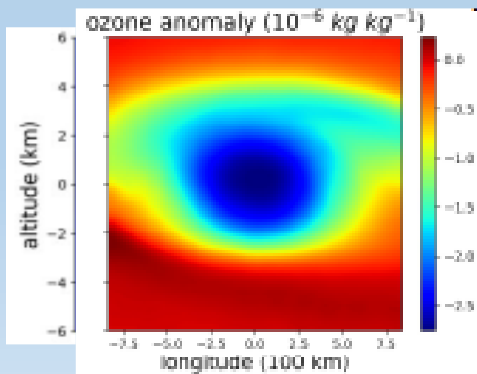
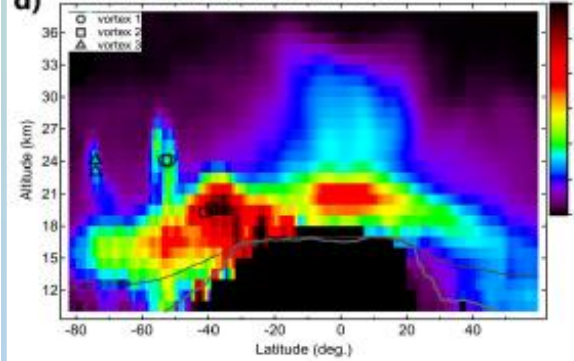
b) MLS water vapour (ppmv), blob sampling



c) MLS carbon monoxide (ppbv), blob sampling



d) OMP2-LP Extinction ratio, 2020 2 1 DOY032



Time-averaged composite section of the ozone anomaly (ECMWF analysis)

-> deep mini ozone hole, depleted by up to 100 DU

Conclusions

- ❖ Global stratospheric aerosol burden largely modulated by volcanic eruptions and by wildfires (which have tended to increase over the recent years) but other particles types have been reported (e.g. organics in the tropical lower stratosphere)
- ❖ Detection of the ATAL aerosol layer from in situ observations confined in the summer monsoon anticyclone confirming the observations by satellite instruments. → mainly small particles
- ❖ The amplitude of the impact of volcanic eruptions critically depend on the latitude and altitude of injection, on the amount of SO₂ and on the timing of the eruption. Monomodal size distributions for moderate volcanic eruptions. Important role of ash particles which perturb the kinetics of aerosol formation and possibly plume dispersion.
- ❖ Volcanic aerosol plumes quite well captured by global models. Uncertainties mainly resulting from differences in resolution and uncertainties in the knowledge of injection parameters.
- ❖ Fires: difficult to make conclusions regarding the generality of smoke vortice structures and what their global impact may be. Compact and vortex-like structures possibly reported after the 2019 eruption of Raikoke (Chouza et al., 2020).
- ❖ However, comparisons between the model and in situ observations show contrasted results and not systematically consistent between model and satellite observations.
→ Difficult to draw robust conclusions if comparisons are not statistically significant.
- ❖ No reference instrument for in situ observations of aerosol concentrations and size distributions. Simultaneous observations of different optical particle counters needed and initiated by HEMERA.
- ❖ OPC observations not sufficient. Need for chemical composition characterization. Ongoing cooperation with NASA-Langley using new balloon-borne aerosols collectors (JP Vernier)