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Evoluzione termofisica dei corpi minori del Sistema Solare

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## **Definition of Small Solar System Bodies**

#### (1) A **PLANET** IS A CELESTIAL BODY THAT:

a) IS IN ORBIT AROUND THE SUN;

b) HAS SUFFICIENT MASS FOR ITS SELF-GRAVITY TO OVERCOME RIGID BODY FORCES SO THAT IT ASSUMES A HYDROSTATIC EQUILIBRIUM (NEARLY ROUND) SHAPE;

HAS CLEARED THE NEIGHBOURHOOD AROUND ITS ORBIT.

(2) A "DWARF PLANET" IS A CELESTIAL BODY THAT:

IS IN ORBIT AROUND THE SUN;

b) HAS SUFFICIENT MASS FOR ITS SELF-GRAVITY TO OVERCOME RIGID BODY FORCES SO THAT IT ASSUMES A HYDROSTATIC EQUILIBRIUM (NEARLY ROUND) SHAPE;

c)HAS NOT CLEARED THE NEIGHBOURHOOD AROUND ITS ORBIT; d)IS NOT A SATELLITE. (3) All other objects, except satellites, orbiting the Sun shall be referred to collectively as "<u>Small Solar</u> System Bodies".

## Why studying small bodies?

## **SMALL BODIES** ARE OBJECTS THAT HAVE PRESERVED INFORMATION ABOUT THE PRIMORDIAL PHASES OF OUR SOLAR SYSTEM.

Information about the **primordial processes** of the **planetary formation**, **water origin** on Earth, **water role** in the thermophysical evolution of planets, role of the **organics**, etc...

#### What is the planetary differentiation?

- DIFFERENTIATION MEANS TO MAKE A
  HOMOGENOUS BODY HETEROGENEOUS.
- IT IS THE PROCESS THAT LEADS TO THE STRATIFICATION OF THE ASTEROID/PLANET BASED ON ITS PHYSICAL AND/OR CHEMICAL PROPERTIES.



#### DIFFERENTIATION

**MINIMIZATION OF GRAVITATIONAL POTENTIAL ENERGY** 

### Main stages of the life of a small body

#### - Accretion

- Sintering
- Core Formation
- Mantle convection
- Cooling phase



## **Minimum Size for Differentiation**

The **minimum size** for the **differentiation** will be set by the rate of **conductive heat loss**. The **characteristic timescale** for the **thermal conduction** into a body of radius R is **R<sup>2</sup>/K**, where K is the thermal diffusivity.

The **minimum size** for the **differentiation** will occur when the heat flux is maximum (i.e.instantaneous accretion). The approximate conductive cooling time for a planetesimal with **instanenous accretion** is given by (<u>Moskovitz & Gaidos 2011</u>):

 $\tau_{cool} = 0.014 Ma \left(\frac{R}{1km}\right)^2$ 

**Planetesimals** with  $\tau_{cool}$  longer than the most important heating timescale ( $\tau_{Al26}$ ) will sustain melting temperatures and differentiate. The minimum size is **18 km**.

## Hydrostatic equilibrium & Differentiation

Can we assume that a **differentiated body** is also in **hydrostatic equilibrium** (HE)?

HE implies that the body has a spherical shape (neglecting the rotation) to minimize gravitational potential. We can not assume that all differentiated bodies achieved HE, even if small asteroids are differentiated, it is unlikely they have sufficient mass to reach spherical form.



#### HYDROSTATIC EQUILIBRIUM



#### **NOT HYDROSTATIC EQUILIBRIUM**







HYDROSTATIC EQUILIBRIUM

#### **Moment of Inertia Factor & Differentiation**

In Planetary Science, momentum of Inertia factor (MoI) is linked to the radial distribution of mass inside a celestial body. For a differentiated planet or satellite, where there is an increase of density with depth, C/MR<sup>2</sup>< 0.4 (which corresponds to the homogeneous sphere).</p>



#### Size & Differentiation



#### Sources of Energy

- The decay of short-lived radioactive elements such as 26AI (primary heat source), 60Fe or of long-lived radioactive elements such as 40K, 232Th, 235U and 238U;
- Accretion process;
- The **differentiation process**, i.e. the separation of the metallic component to the silicatic one;
- Tidal heating:
- The impacts with other bodies;
- Electrical induction by an early T-Tauri like solar wind;



#### **Energy Source Requirements**

I) **POWERFUL**, since small bodies have a high surface/volume ratio;

II) FAST, as suggested by the measurements on Hf – W ratio in iron meteorites (Kleine 2002);

II) UNIFORMLY DISTRIBUTED in our galaxy

## **Accretional Heating**

The fraction (h) of heat retained during accretion is very poorly known. In terms of h, we can approximate the accretional temperature T<sub>A</sub> at radius r within an accumulating planetesimal as (<u>Schubert et al. 1986</u>):

$$T_A(r) = \frac{hGM(r)}{c_v r} \left[ 1 + \frac{rv^2}{2GM(r)} \right] + T_E$$

- If accretion is slow, heat will be radiated into space faster than it is delivered and the body will not heat up. If accretion is fast, heat will be stored and the temperature of the body will increase. Depending on the amount stored heat and on the size of the body, the accretion profile can rise the melting temperature of iron and possibly of silicate.
- In the optimistic case in which the **50%** of the **energy** is **converted** in **heat** and stored in the asteroid, i.e. h = 0.5, the **temperature increase** is only about **50K**.



For the **Earth**, **accretion** is the main source of energy:  $\Delta T \ge 1000 \text{ K}$ (Solomon 1989)

- C<sub>v</sub> = specific heat;
- $T_{e}$  = ambient temperature;

V<sup>2</sup>/2 = approach kinetic energy per unit of mass; M(r) = mass of the accreting asteroid internal to r; h = efficiency.

#### **Short-lived Radionuclides**

Parent nuclide	<sup>26</sup> Al	<sup>60</sup> Fe	<sup>53</sup> Mn
Daughter nuclide	<sup>26</sup> Mg	<sup>60</sup> Ni	<sup>53</sup> Cr
Initial isotopic abundance	$^{26}$ Al/ <sup>27</sup> Al = 5 × 10 <sup>-5</sup>	${}^{60}$ Fe/ ${}^{56}$ Fe = 0.1–1 × 10 <sup>-6</sup>	$^{53}$ Mn/ $^{55}$ Mn = 1–4 × 10 <sup>-5</sup>
Half-life $\lambda$ (Myr)	0.716-0.73	1.5	3.7
Specific heat production (W/kg)	0.146	0.068-0.074	0.027

(from <u>Castillo-Rogez et al. 2007</u>)





## Long-lived Radionuclides

The most important long-lived radionuclides are 238U, 235U, 232Th and 40K: decay information about these elements are reported in the following table (<u>Castillo-Rogez et al. 2009</u>) :

Element	Potassium	Thorium	Uranium	
Isotope	<sup>40</sup> K	<sup>232</sup> Th	<sup>235</sup> U	<sup>238</sup> U
Isotopic abundance (wt%)	0.01176	100.00	0.71	99.28
Decay constant $(yr^{-1})$	$5.54 \times 10^{-10}$	$4.95 \times 10^{-11}$	$9.85 \times 10^{-10}$	$1.551 \times 10^{-10}$
Half-life $\lambda$ (Myr)	1277	14,010-14,050	703.81	4468
Specific heat production (W/kg of elements)	$29.17 \times 10^{-6}$	$26.38 \times 10^{-6}$	$568.7 \times 10^{-6}$	$94.65 \times 10^{-6}$

These long-lived radionuclides are still present in the Earth and other planetary bodies, although in lower abundances with respect to the Solar System formation. This allows for sustained long-term heating during planetary evolution.



#### Magma Ocean & Core Formation



If **T** > 1400 K (<u>Neumann et al. 2013</u>) -> Magma Ocean (<u>Righter & Drake 1997</u>) in which iron drops migrates at the bottom creating a core (<u>Stevenson 1990</u>).

The **velocity** of **iron drops** is directly proportional to the density difference with the silicate matrix and to the radius<sup>2</sup> and inverse proportional to the viscosity (**Stoke law**).

**Crystallization sequence** in the magma ocean follows the **Bowen's series**.

#### Some Results: Vesta



**26AI affinity** with **silicates** leads to the formation of a peculiar **temperature** profile, with the **peak** in the **mantle** (<u>Ghosh & McSween 1998</u>, Formisano et al. 2013A).



(from Ghosh & McSween 1998)

#### Some Results: Lutetia







50

50

(from Formisano et al. 2013B)

#### Heat transfer mechanisms

- CONDUCTION (the most diffused);
  CONVECTION
- RADIATION



#### Most materials...

- Get less dense when heated
- More dense when cooled
- Hot stuff buoyant relative to surroundings: rises
   Cold stuff relatively heavy: falls

#### **Thermal Convection**



The **Rayleigh number** characterizes the competition between forcing by thermal bouyancy and damping by viscosity and thermal diffusion

Ra must exceed a certain value called the critical Rayleigh number  $(Ra_c)$  for convection to occur. For  $Ra < Ra_c \rightarrow$  the layer is stable and trasports heat by conduction For  $Ra > Ra_c \rightarrow$  the layer will be convectively unstable and trasports heat more rapidly via convection.

# Thermal Boundary Layers & Nusselt Number

With very large Ra and vigorous mixing most of the layer is largely uniform and isothermal. Most of the fluid is at the mean temperature between the two boundary temperatures.

However the **temperature** still must **drop** from the well-mixed warm interior to the cold temperature at the top, and to the hotter temperature at the bottom. The **narrow regions** accommodating these jumps in temperature are called **thermal boundary layers**.

Nu = Nusselt number= Heat transfer across layer in convective regime/ Heat transfer across layer by conduction

 $Nu = q_{convection}/k\Delta T/d$ 



#### (from Turcotte & Schubert 2002)

temperature

#### **Thermal Convection "typologies"**



If the **viscosity** is high (>10<sup>19</sup> Pa s, as in the asthenosphere – <u>Schubert</u> <u>1979</u>) -> "Sub-Solidus Convection" If the **viscosity** is low (**few Pa s**) -> Core convection

#### Sub-solidus Convection: Ceres' crust



(from Formisano et al. 2019, JGR, submitted)

## Core Dynamo as a consequence of the Differentiation

- Three conditions are required to drive a dynamo on a growing planet (Monteux et al. 2011) and so to have a magnetic field:
- I) The heat flow out of the core must exceed the adiabatic value such that convection can occur (<u>Stevenson et al. 1983</u>).
  - 2) The **ratio** of the **rate** at which **gravitational potential energy** is **released** by **convection** to the **rate** of **ohmic dissipation** must exceed a critical value (<u>Buffett 2002</u>).
- 3) The structure of the convective motions carrying magnetic field lines must be sufficiently complicated to favor self-sustaining dynamo action, i.e. if magnetic Reynolds number overcomes a critical value (typically in the range 10-100 (Monteux et al. 2011 and reference therin)).

**Typical timescale** of **core dynamo** is < **100 Ma** (<u>Elkins-Tanton et al. 2011</u>)

#### **Core Convection: Vesta & Psyche**

- **Vesta** probably had a magnetic field as suggested by the studies on the residual magnetization on an eucritic meteorite (Fu et al. 2012) and suggested by theoretical models (Formisano et al. 2016);
  - **Psyche** could have been characterized by a core dynamo (<u>Bryson et al. 2015</u>), appearing to be an **exposed metallic core** or a **fragment** of a **metallic core** from a **larger differentiated parent body** some 500 kilometers in diameter. Its origin is still matter of debate.



### Why studying magnetic fields?

- Information about:
- The **physics** of the **interior** of the celestial bodies;
- ii. Past geological activities;
- Mathematically interesting since it is a problem of complex system

ALL THE **EFFECTS** OF THE **PAST MAGNETIC FIELDS** ARE RECORDED IN THE ROCK IF THE LOCAL TEMPERATURE DOES NOT OVERCOME THE **CURIE TEMPERATURE**.

#### Conclusions

- Small bodies are important to understand the main stages of the evolution of our Solar System;
- Numerical models offer a theoretical support to the observations, in order to select the most plausible physical scenarios:
- New models (and new computers) are required to simulate more and more complicated physical situations, in particular turbulent convection and core dynamos generation.